

Chapter 5

Transforming Geocentric Coordinates to Official Coordinates

Module MG E-01: Coordinate Systems

Dr.-Ing. Christoph Holst

&

► Dr.-Ing. Tomislav Medic

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Institute of Geodesy and
Geoinformation

- 1. Geocentric coordinate systems**
- 2. Need for official coordinates**
- 3. Official coordinates in Germany**
- 4. Transformation of GNSS coordinates to official coordinates**
- 5. Transformation of TLS coordinates to official coordinates**

- Global geocentric system (ECEF, Earth Centered Earth Fixed)
- Used to describe the position on Earth
- Origin: Center of Mass of the Earth
- Z-axis through rotation axis
- X-axis is intersection between Greenwich meridian and equatorial plane
- Y-axis completes right handed system

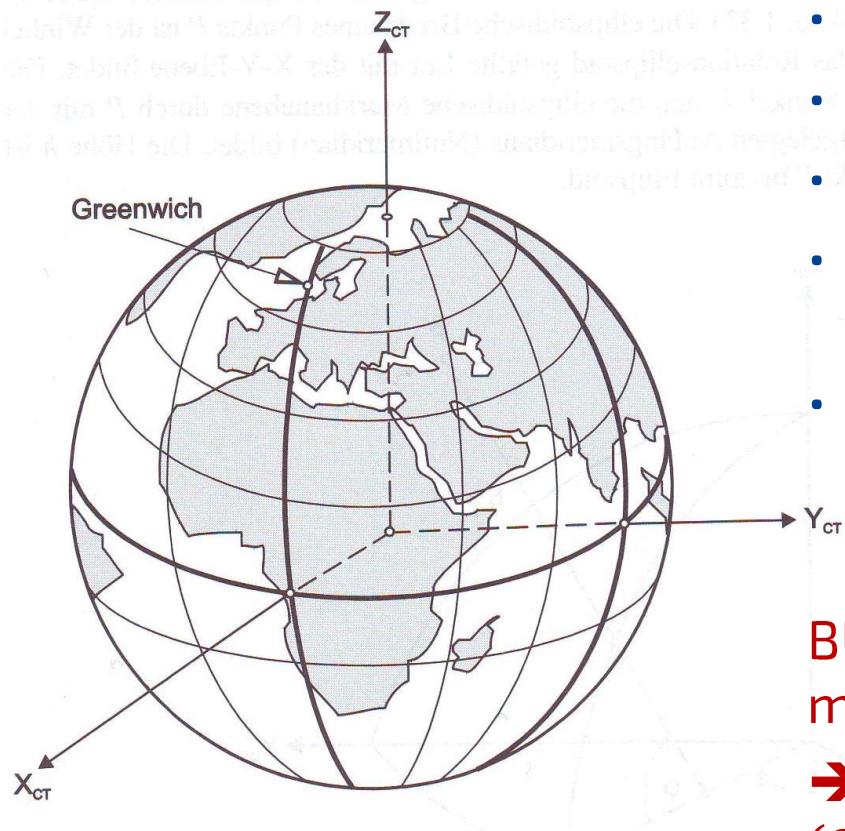
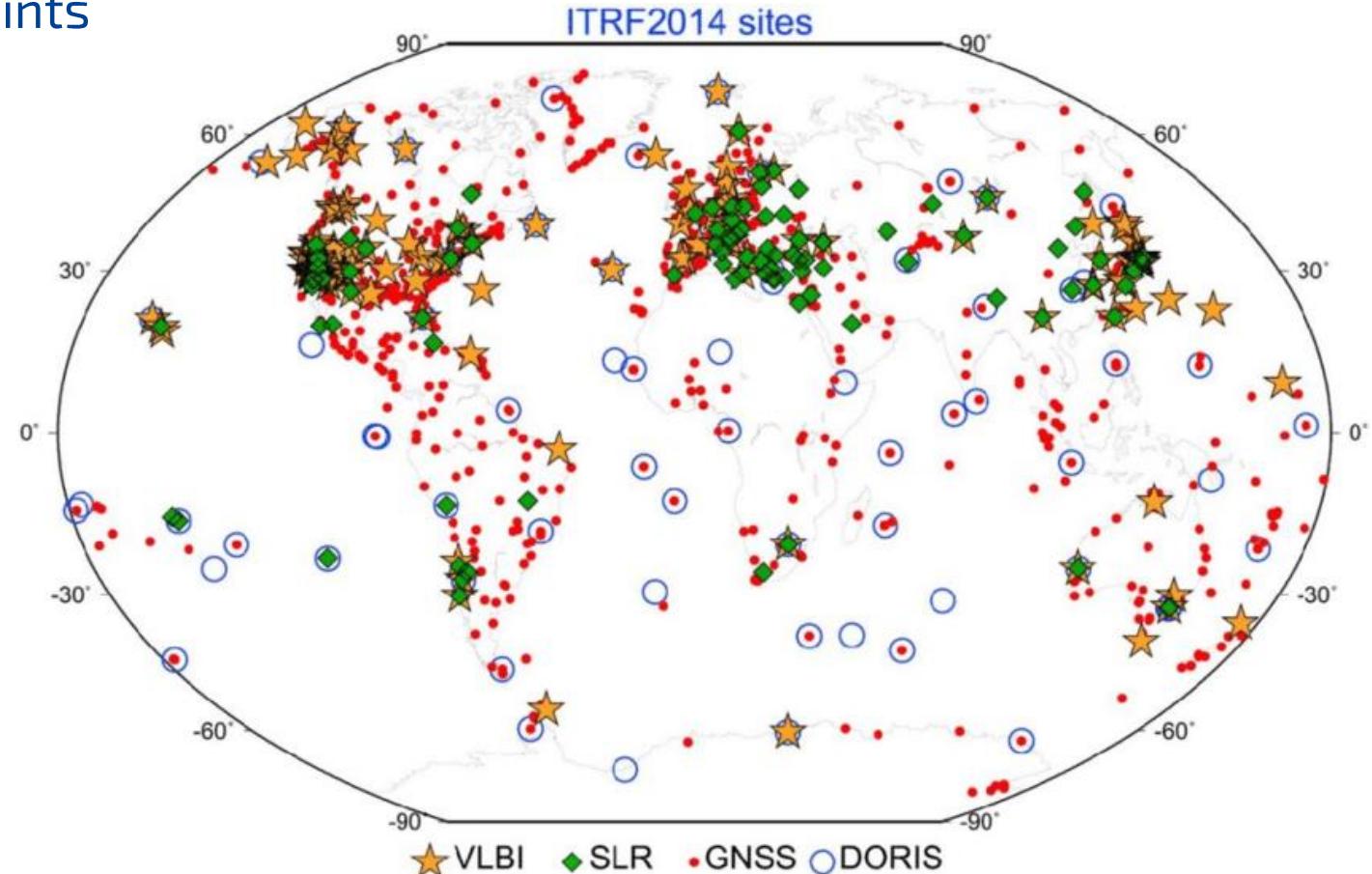


Fig.: Global geozentrisches Koordinatensystem
Source: Bauer, 2011

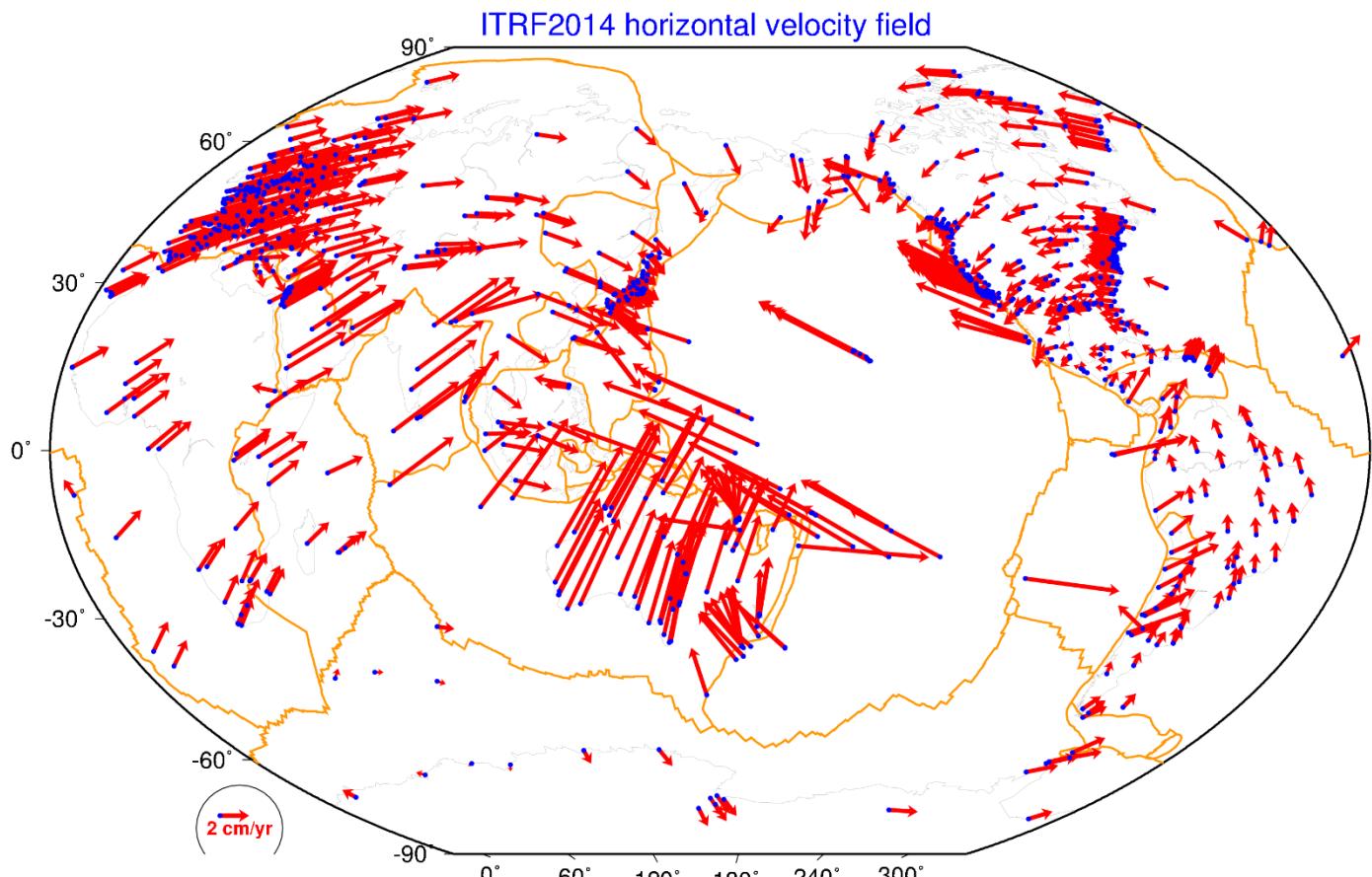
BUT: rotation axis is not constant (polar motion)
→ z-Axis is going through the CTP
(Conventional Terrestrial Pole, the average of the Earth pole between 1900 and 1905)

- Reference Frames are realizations of Reference Systems
- Axis and origin are not physically accessible AND change in time
=> the system is implemented by providing the coordinates at a number of dedicated points



Source: gpsmet.agt.bme.hu

- Because of the motion of the Earth crust, the points have to be updated regularly and a velocity of the points is also given



Source: gpsmet.agt.bme.hu

- **ITRFyy**

- International Terrestrial Reference Frame, calculated in the year yy
- Maintained by the Scientific Community
- Current version is ITRF2014
- Accuracy 5-15mm and 2-3mm/year

- **ETRS89**

- European Terrestrial Reference Frame
- Frozen ITRF89
- **Definition of official coordinates in Germany**

- **WGS84**

- World Geodetic System
- Developed by Department of Defense (US), Military System
- Also specifies a reference ellipsoid, a set of fundamental constants and an Earth gravity model
- Adapted to the ITRF, differences are in the order of decimeters
- **Definition of GPS Satellites and corresponding user coordinates**

• ITRFyy

- International Terrestrial Reference Frame, calculated

• WGS84

- World Geodetic System
- Developed by Department

- Similarity Transformation between coordinate systems:

$$\begin{bmatrix} X' \\ Y' \\ Z' \end{bmatrix} = \lambda \underbrace{\mathbf{R}_3(\varepsilon_z) \mathbf{R}_2(\varepsilon_y) \mathbf{R}_1(\varepsilon_x)}_{\mathbf{R}} \begin{bmatrix} X \\ Y \\ Z \end{bmatrix} + \begin{bmatrix} t_x \\ t_y \\ t_z \end{bmatrix}$$

• ETI

Reference Frame

- Frozen ITRF89
- Definition of official coordinates in Germany

- Since rotations quite small, approximation is valid:

$$\begin{bmatrix} X' \\ Y' \\ Z' \end{bmatrix} = \begin{bmatrix} X \\ Y \\ Z \end{bmatrix} + \begin{bmatrix} t_x \\ t_y \\ t_z \end{bmatrix}$$

decimeters

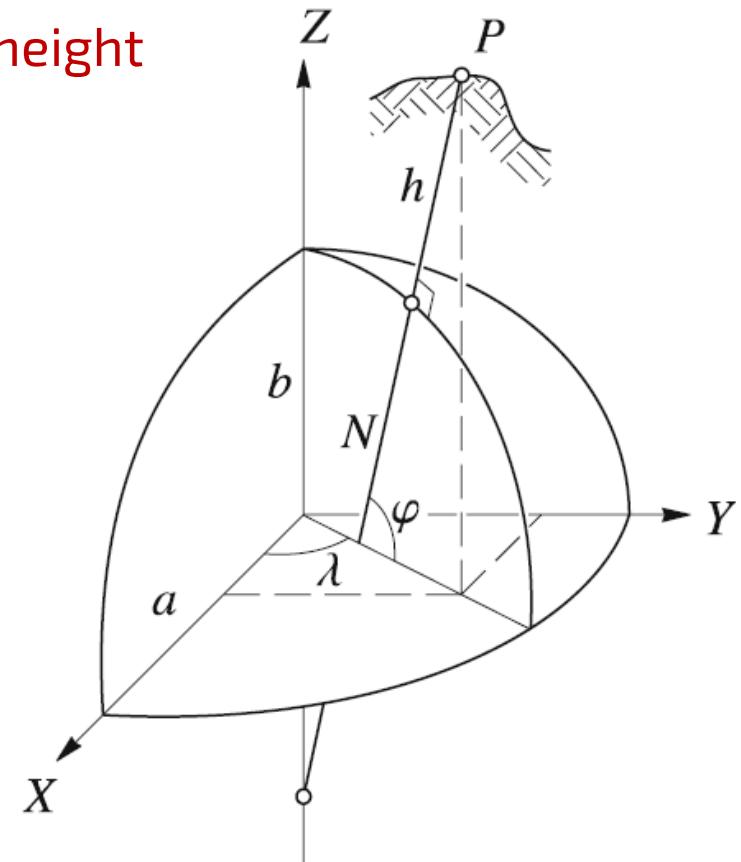
- Definition of GPS Satellites and corresponding user coordinates

- Ellipsoidal coordinates are easier to interpret on earth
- Ellipsoidal coordinates need the definition of an reference ellipsoid (e.g. GRS80, WGS84, Bessel)
- λ ... longitude, φ ... latitude, h ... ellipsoidal height
- $\lambda, \varphi, h \Rightarrow X, Y, Z$

$$\begin{bmatrix} X \\ Y \\ Z \end{bmatrix} = \begin{bmatrix} (N(\varphi) + h) \cdot \cos \varphi \cdot \cos \lambda \\ (N(\varphi) + h) \cdot \cos \varphi \cdot \sin \lambda \\ \left(\frac{b^2}{a^2} N(\varphi) + h \right) \cdot \sin \varphi \end{bmatrix}$$

$$N(\varphi) = \frac{a^2}{\sqrt{a^2 \cdot \cos^2 \varphi + b^2 \cdot \sin^2 \varphi}}$$

... transverse radius of curvature

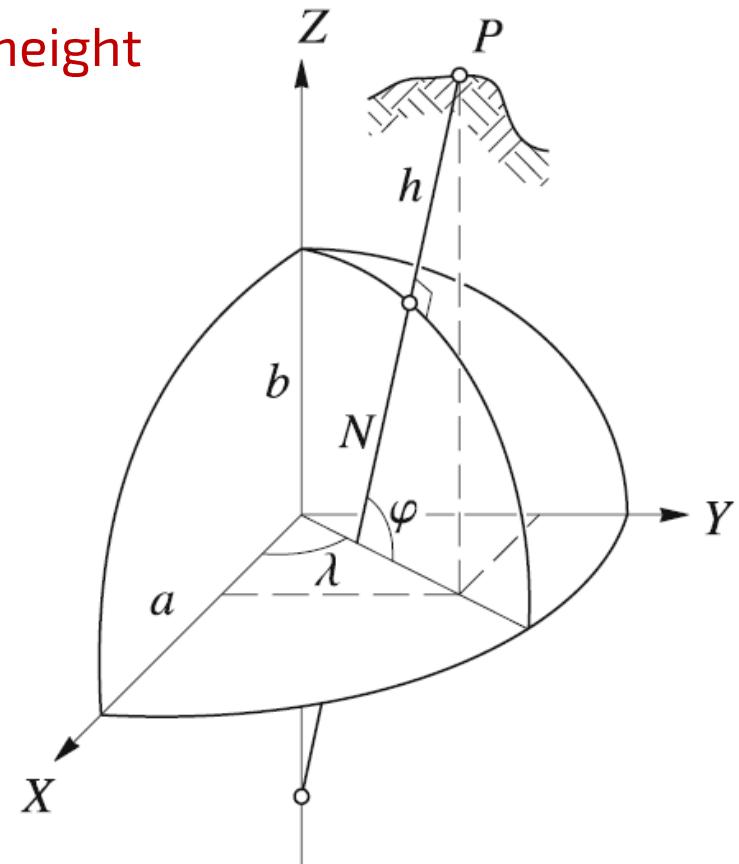


Hofmann-Wellenhof et al. (2008), Fig.8.1

- Ellipsoidal coordinates are easier to interpret on earth
- Ellipsoidal coordinates need the definition of an reference ellipsoid (e.g. GRS80, WGS84, Bessel)
- λ ... longitude, φ ... latitude, h ... ellipsoidal height
- $X, Y, Z \Rightarrow \lambda, \varphi, h$

$$\begin{bmatrix} \varphi \\ \lambda \\ h \end{bmatrix} = \begin{bmatrix} \tan^{-1} \left(\frac{Z}{\sqrt{X^2 + Y^2}} \cdot \left(1 - \frac{a^2 - b^2}{a^2} \cdot \frac{N(\varphi)}{N(\varphi) + h} \right)^{-1} \right) \\ \tan^{-1} \left(\frac{Y}{X} \right) \\ \frac{\sqrt{X^2 + Y^2}}{\cos \varphi} - N(\varphi) \end{bmatrix}$$

- Iterate: $N(\varphi) = \frac{a^2}{\sqrt{a^2 \cdot \cos^2 \varphi + b^2 \cdot \sin^2 \varphi}}$

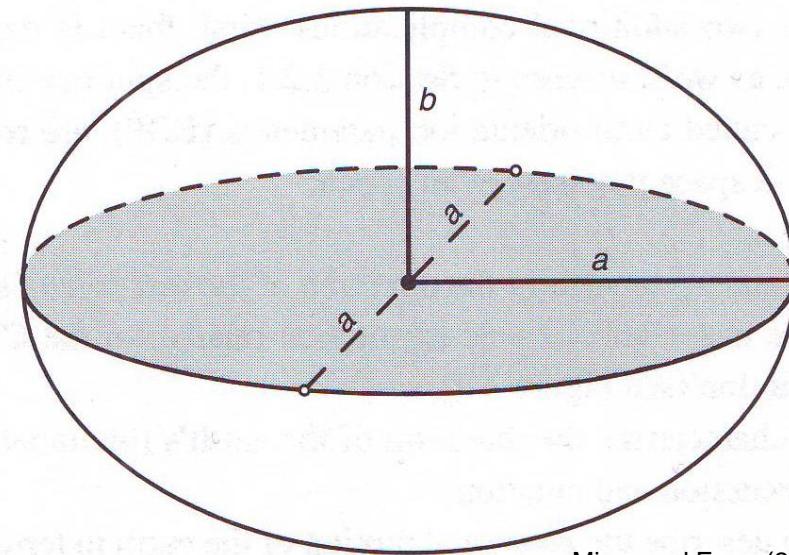


Hofmann-Wellenhof et al. (2008), Fig.8.1

- A reference ellipsoid has to be defined in combination with a reference frame

- $e^2 = \frac{a^2 - b^2}{a^2}$; $f = \frac{a - b}{a}$

- e... eccentricity
- f... flattening
- a... semi-major axis
- b... semi-minor axis



Misra und Enge (2001), Fig. 3.3

	a	1/f
WGS84	6378137.000	298.25722356300
GRS80	6378137.000	298.25722210088
Bessel	6377387.155	299.15281285000
PZ90.11	6378136.000	298.25783930300

- Each coordinate is related to
 - Geodetic datum: ITRFyy, ETRS89, WGS84, ...
 - Reference surface: none = 3d Cartesian, Ellipsoid, Sphere, Plane
- GPS coordinates are given with geodetic datum WGS84 as 3D Cartesian coordinates (= no reference surface)
 - X, Y, Z
- Often, they are transformed to ellipsoidal coordinates for better interpretation (= Ellipsoid reference surface)
 - λ ... longitude, φ ... latitude, h ... ellipsoidal height

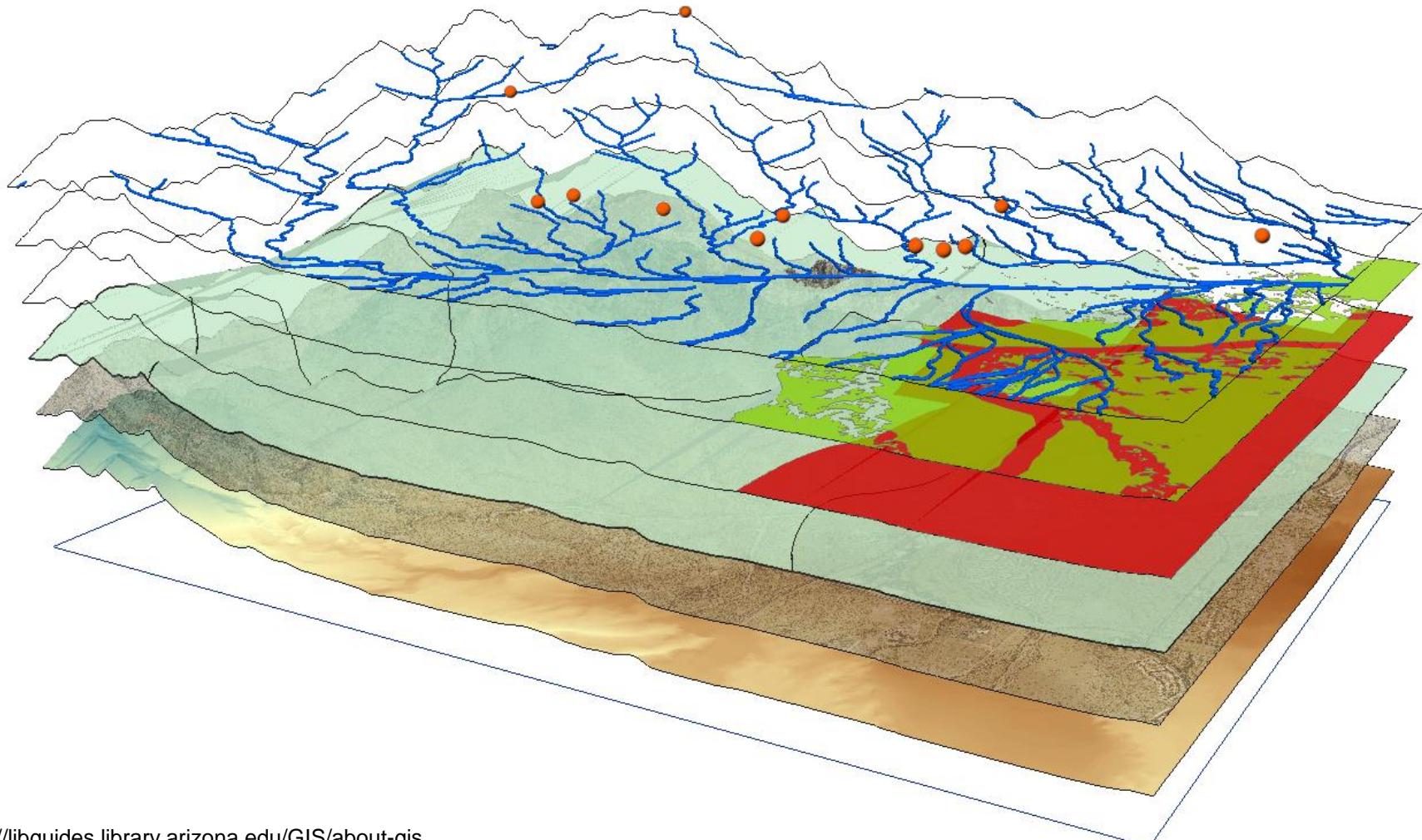
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- Cadaster: only positional coordinates (2D), metric, stable



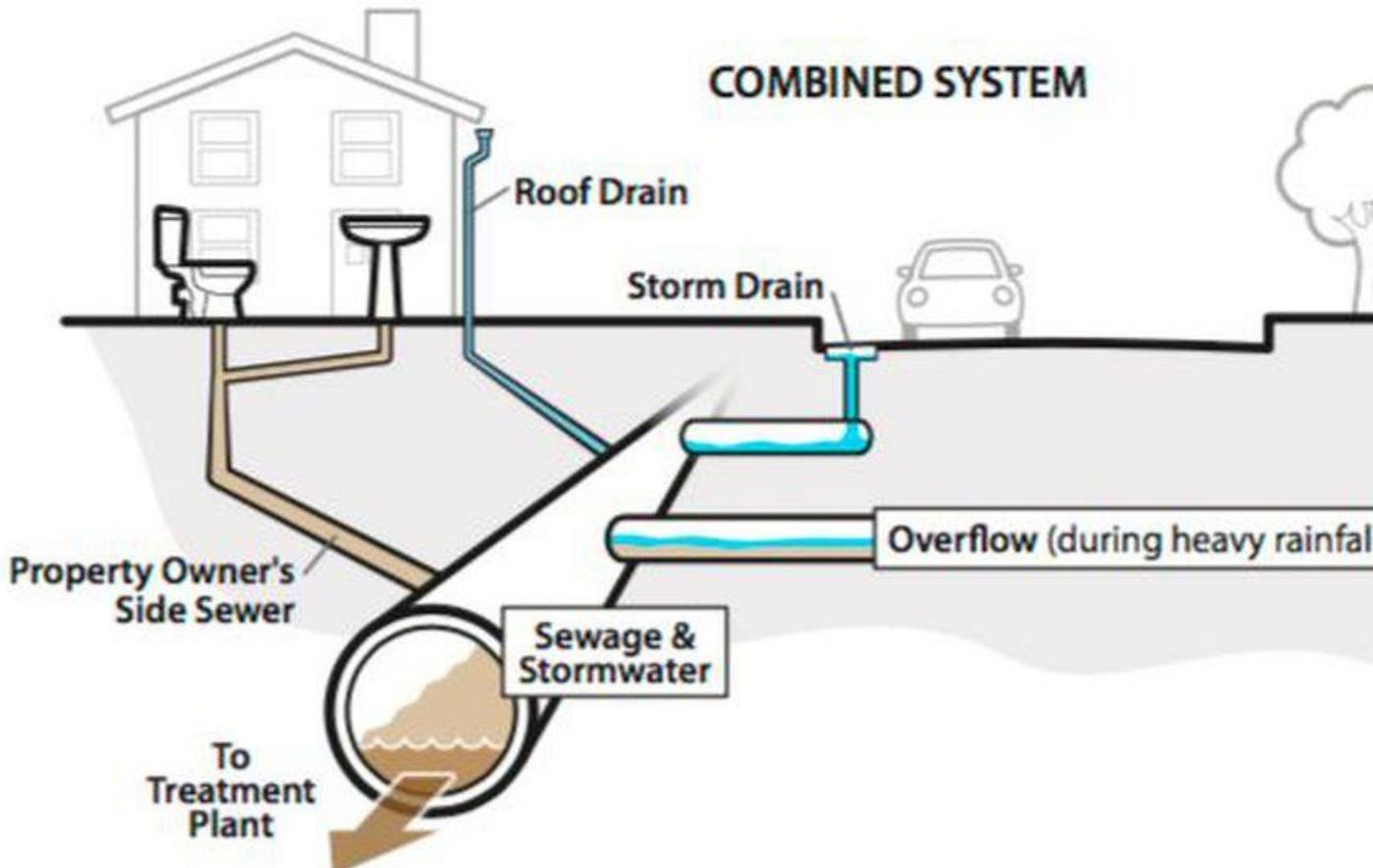
Source: LAiV-MV

- Positional coordinates (+ height), metric, stable



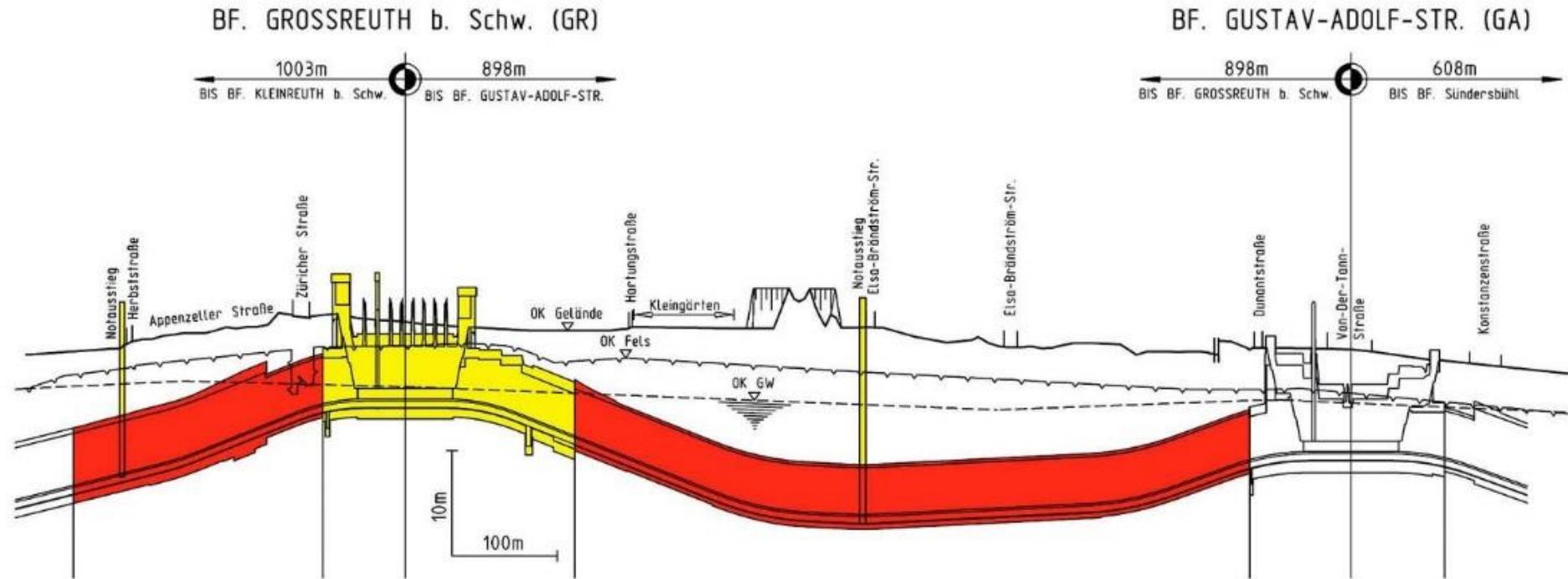
<http://libguides.library.arizona.edu/GIS/about-gis>

- Heights: physical interpretable => no water flows between two points with the same height



Source: City of Seattle

- Positional coordinates and height, metric, stable
- Heights: physical interpretable => no water flows between two points with the same height

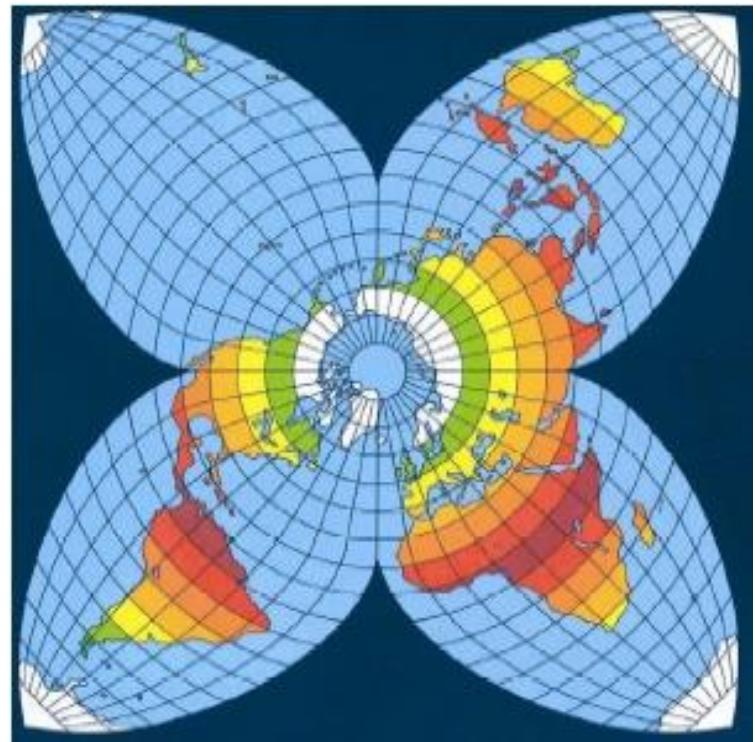


Source: City of Münster

- Official coordinates need to be **separated by position (2D) and height**
- Position needs to be **metric** => no longitude and latitude
- Height needs to be **physical** => no ellipsoidal height
- Official coordinates need to be **stable**, i.e., no coordinate change due to global earth processes => no ITRF / WGS84 for Europe

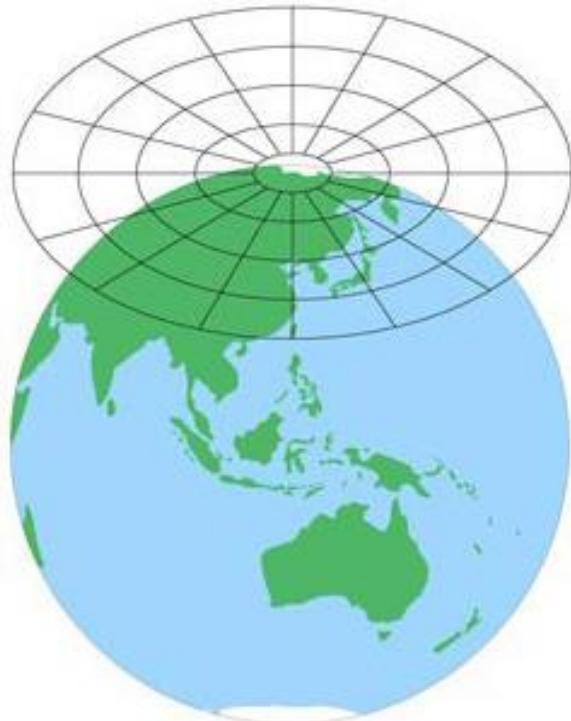
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How to flatten the earth?

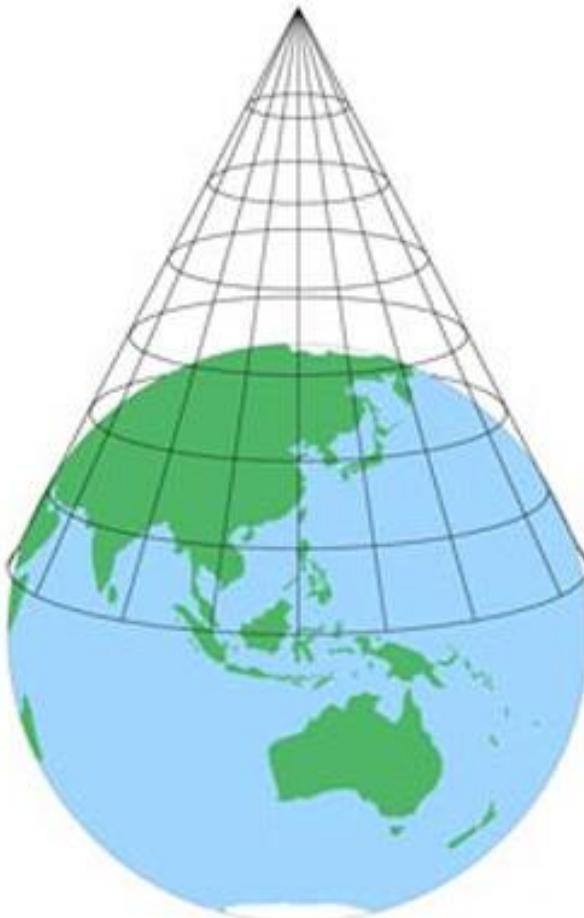


Projection of 3D coordinates in 2D plane =>
Not applicable without distortion

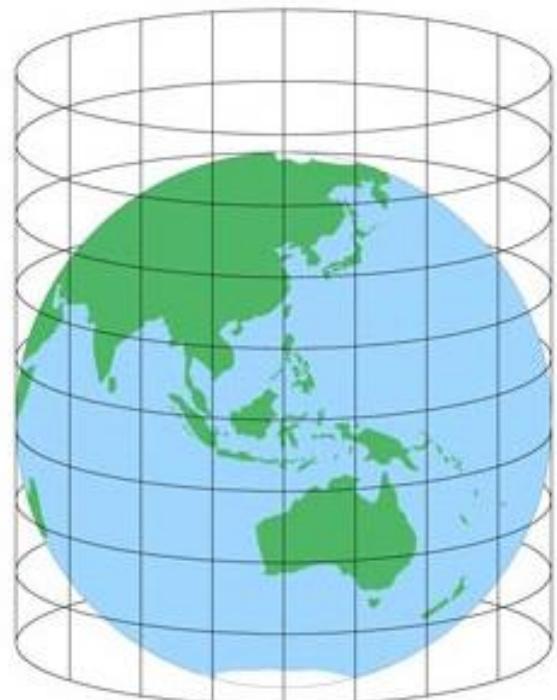
azimuthal



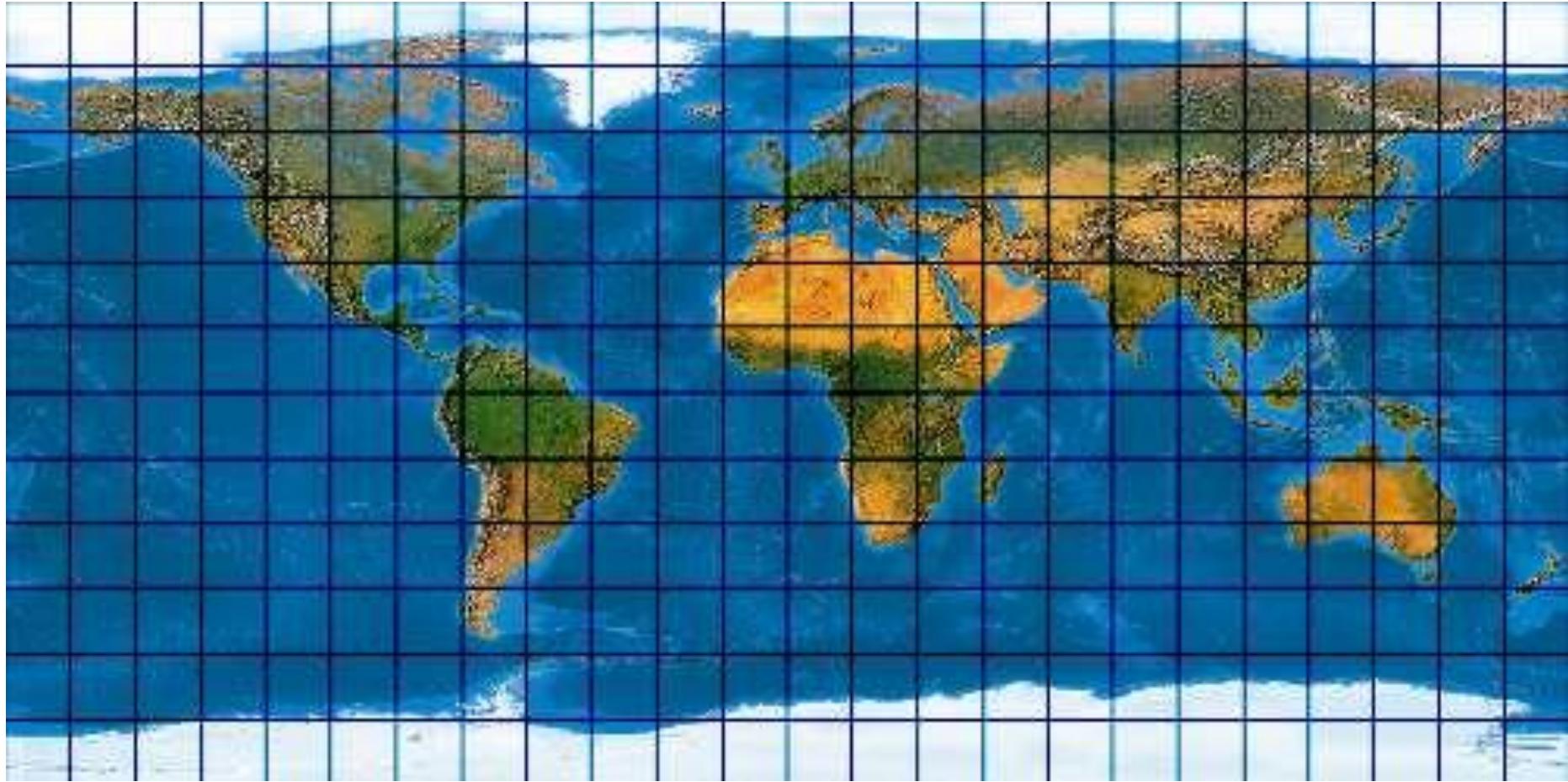
conical



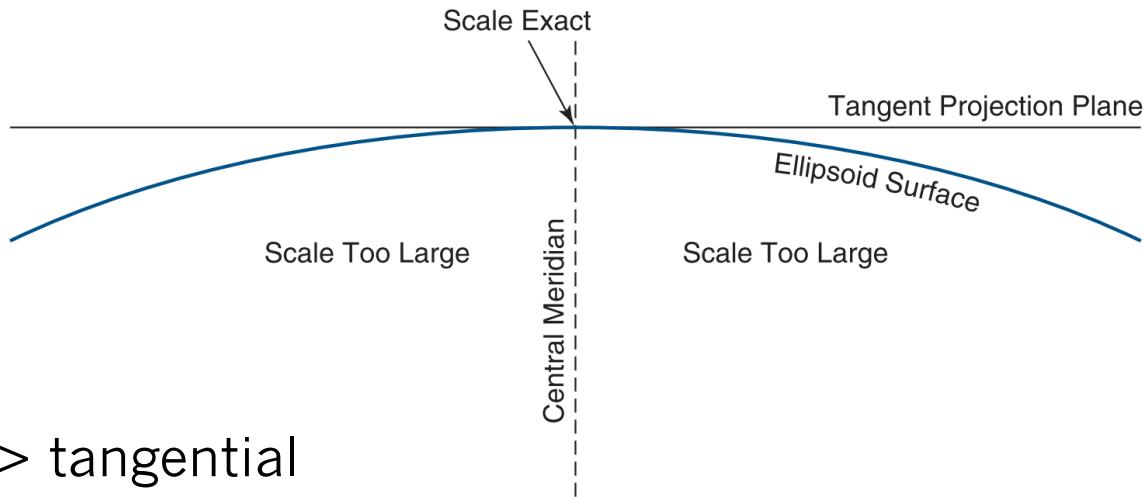
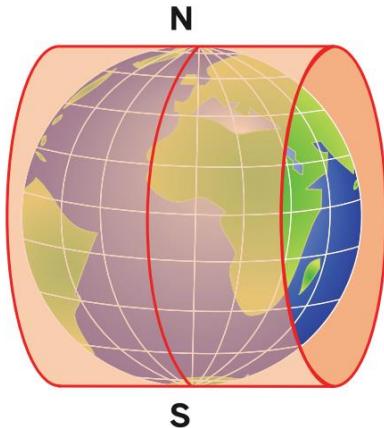
cylindrical



- Rectangular, planar coordinate system
- East axis, north axis

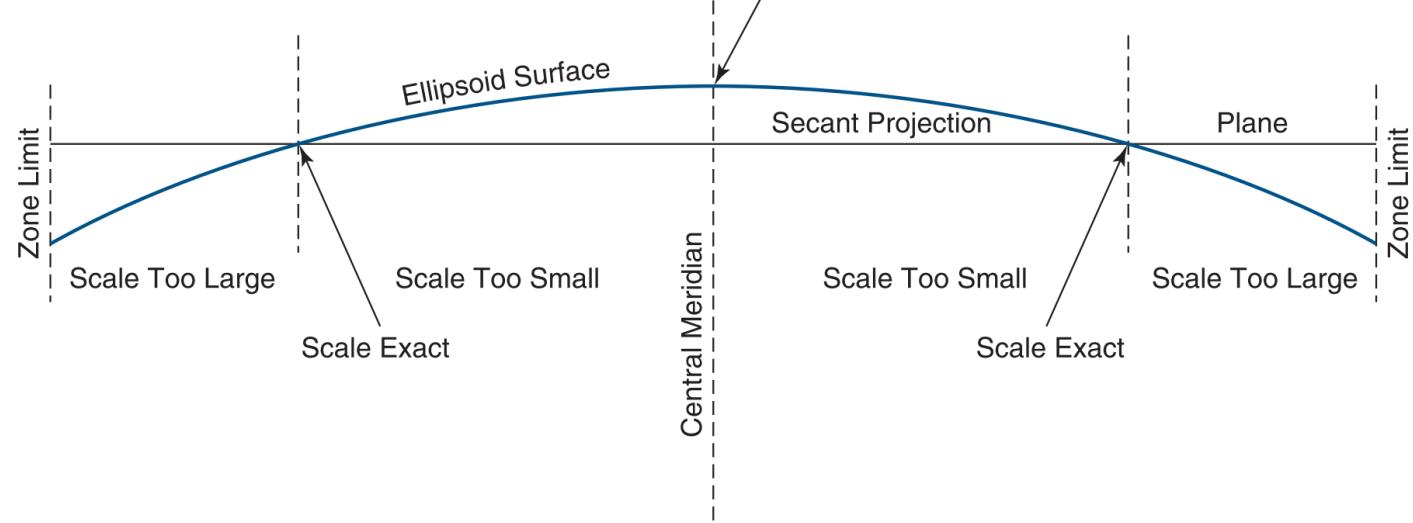
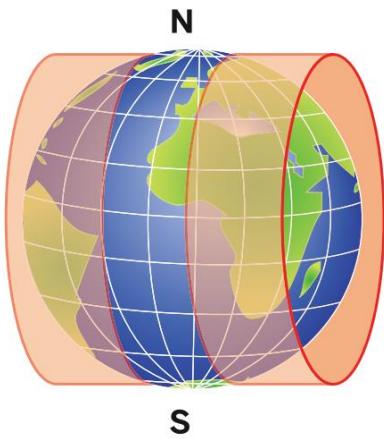


Realization of map projection

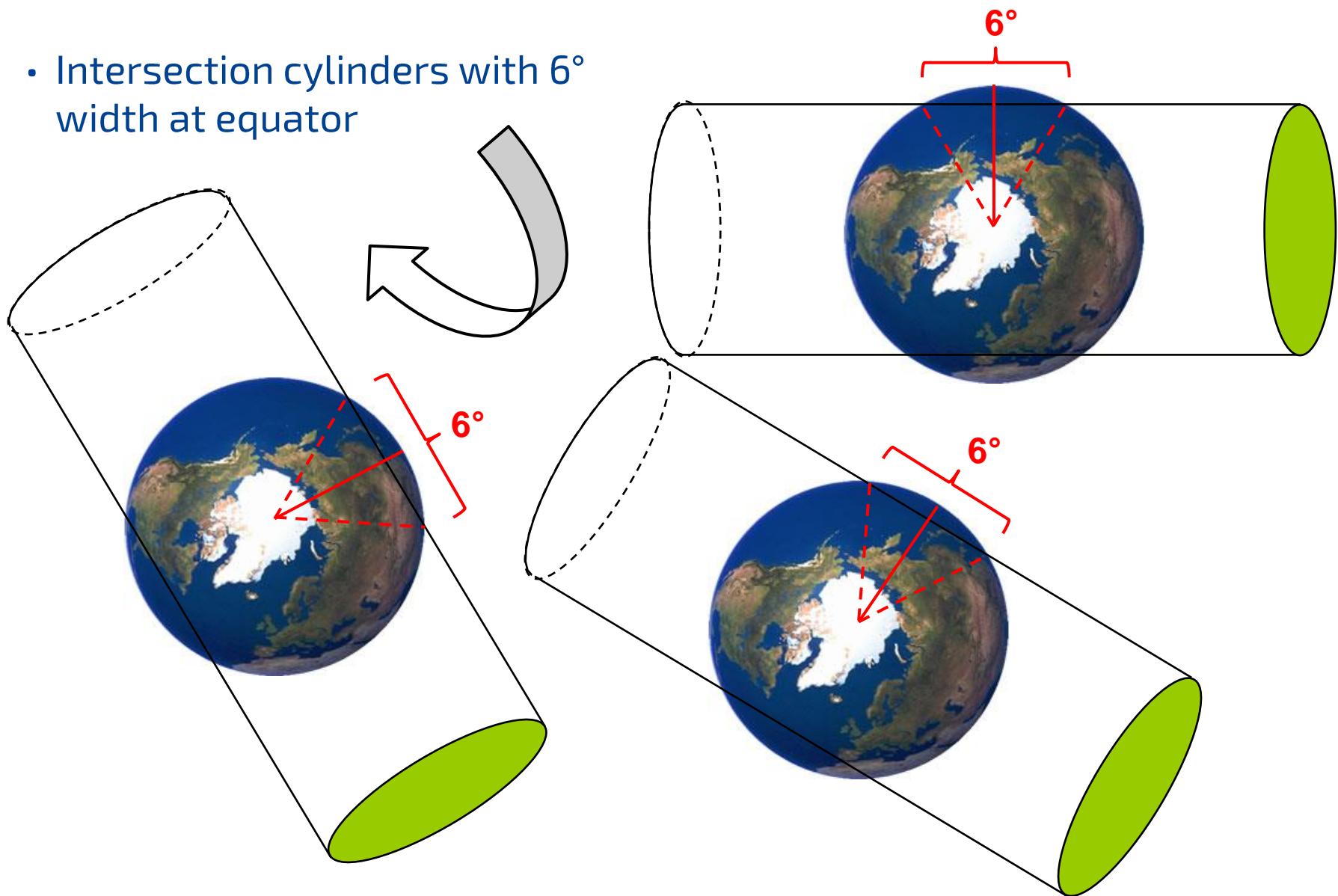


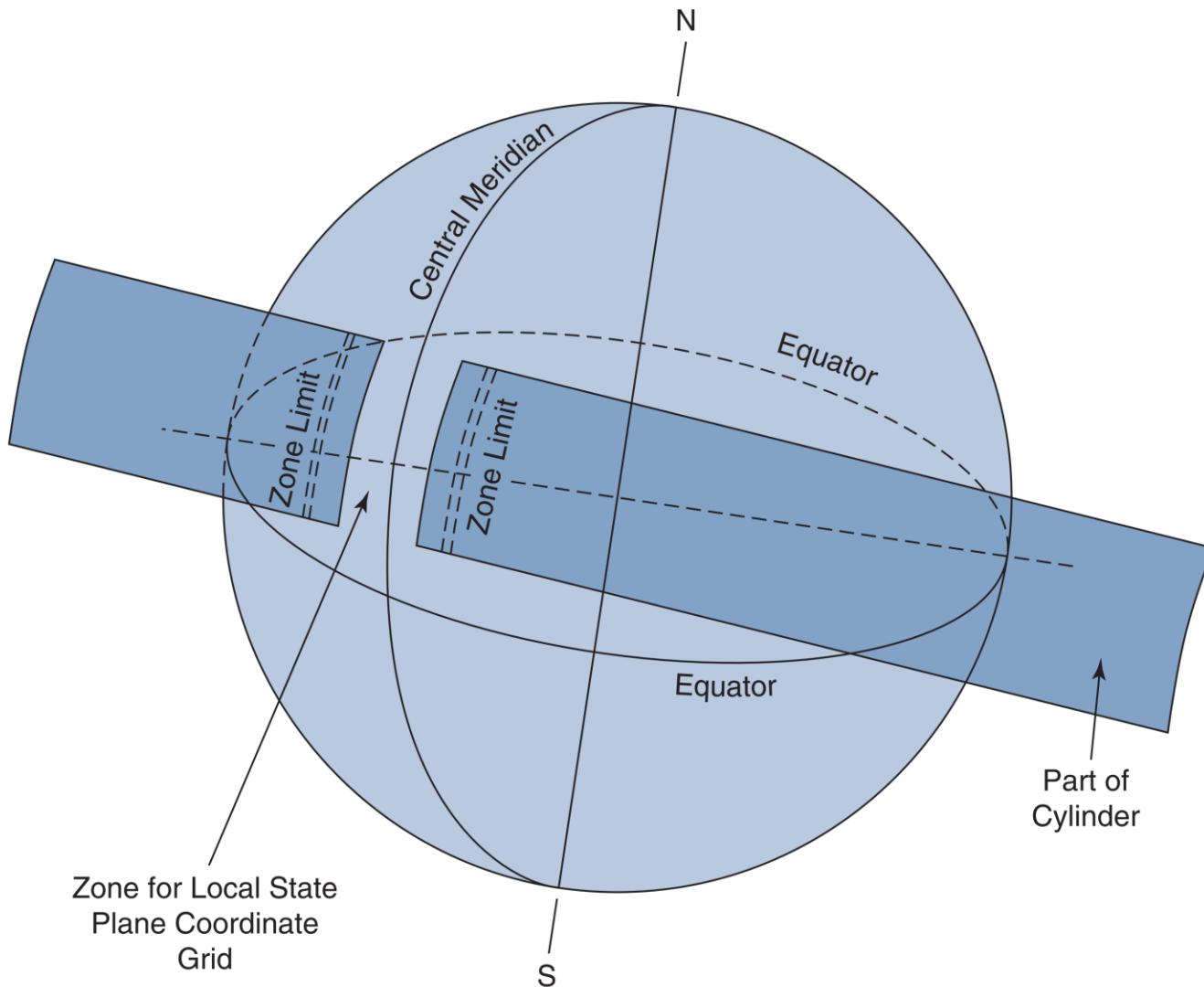
Above: Gauß-Krüger => tangential

Below: **Universal Transversal Mercator (UTM)** => intersectional

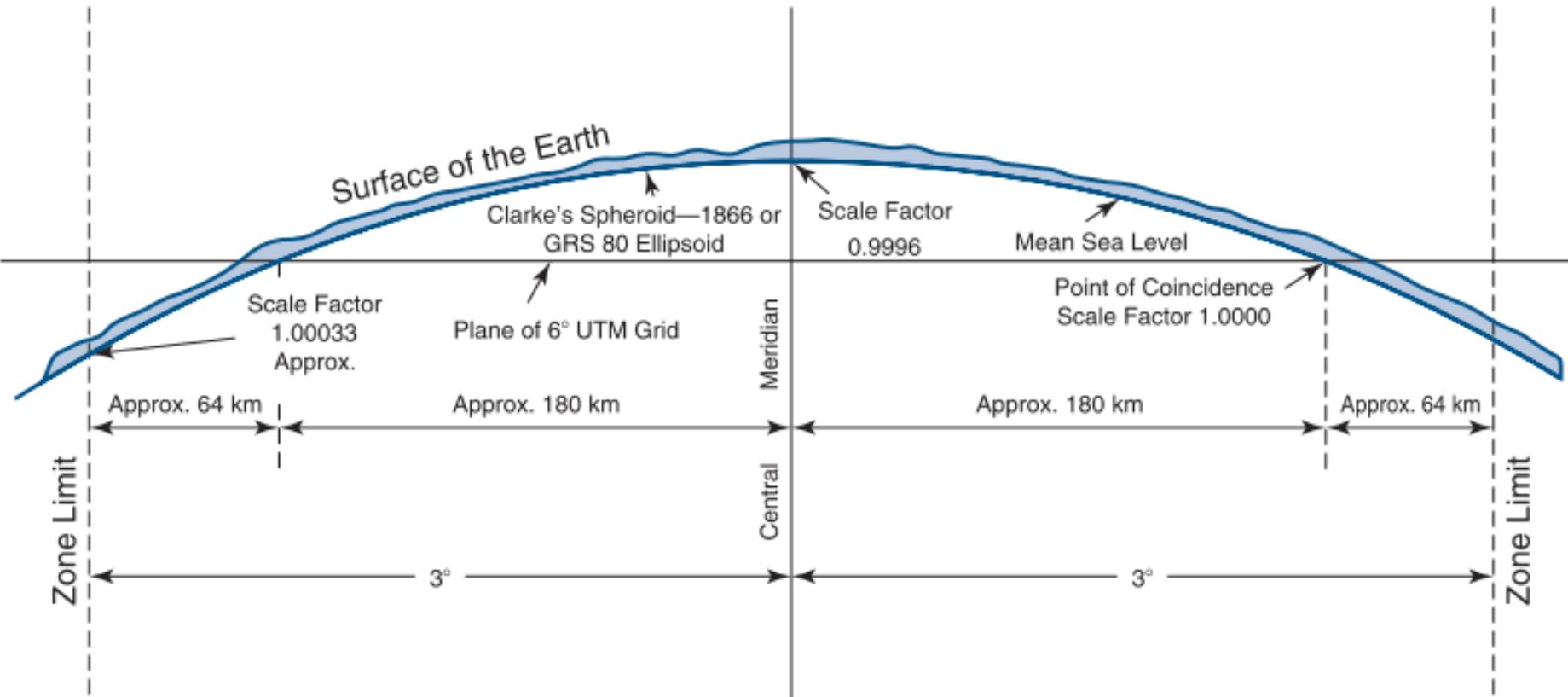


- Intersection cylinders with 6° width at equator



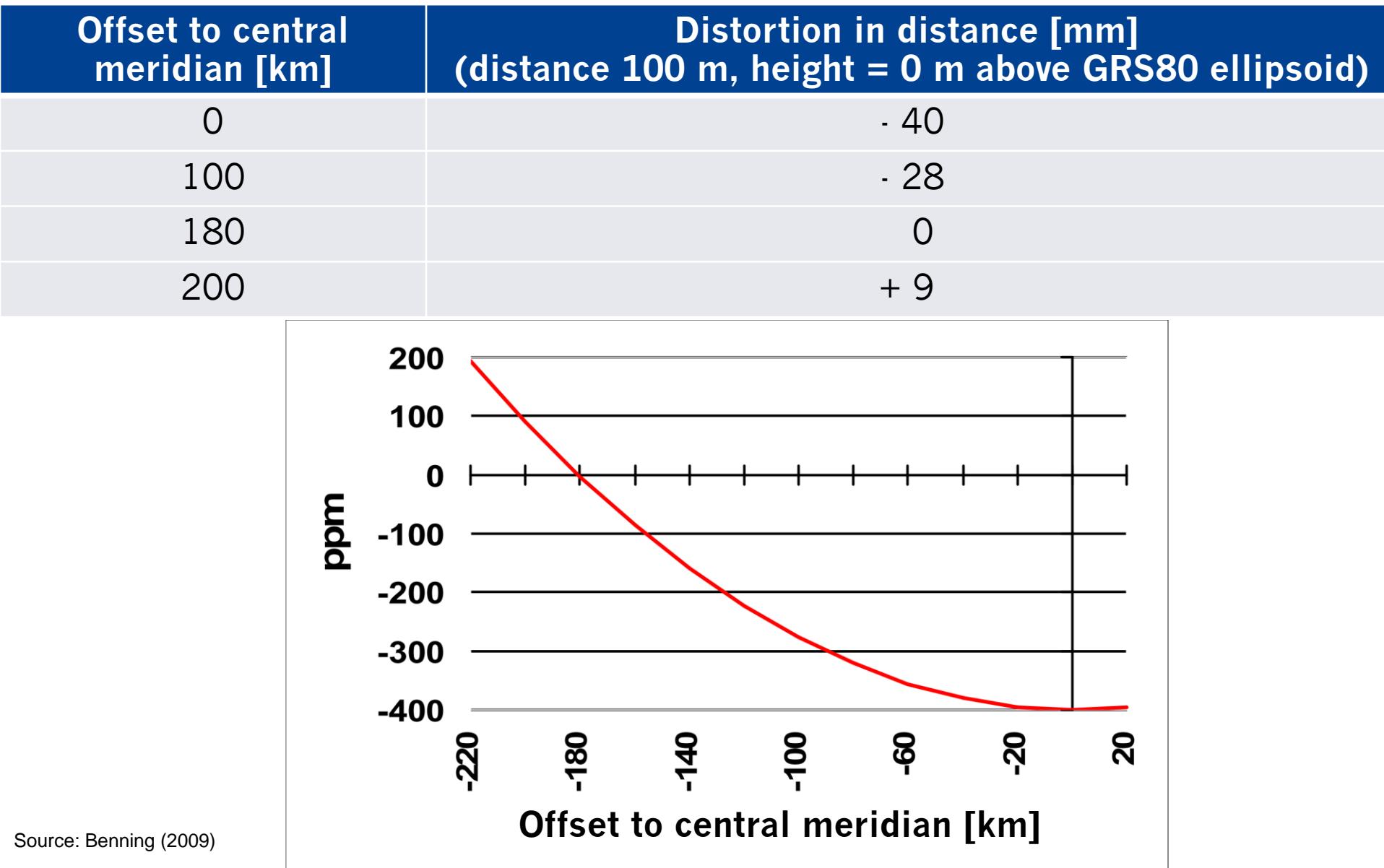


Source:
Kavanagh
(2014)



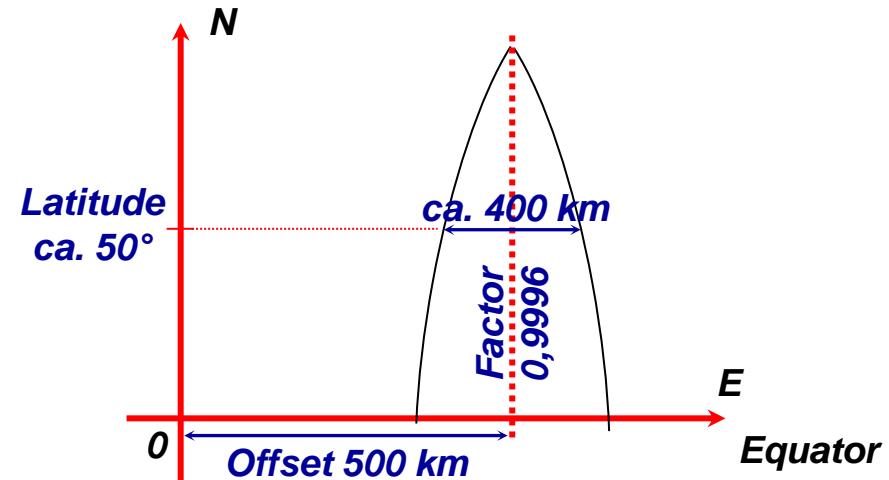
Source:
Kavanagh
(2014)

Distortion due to projection



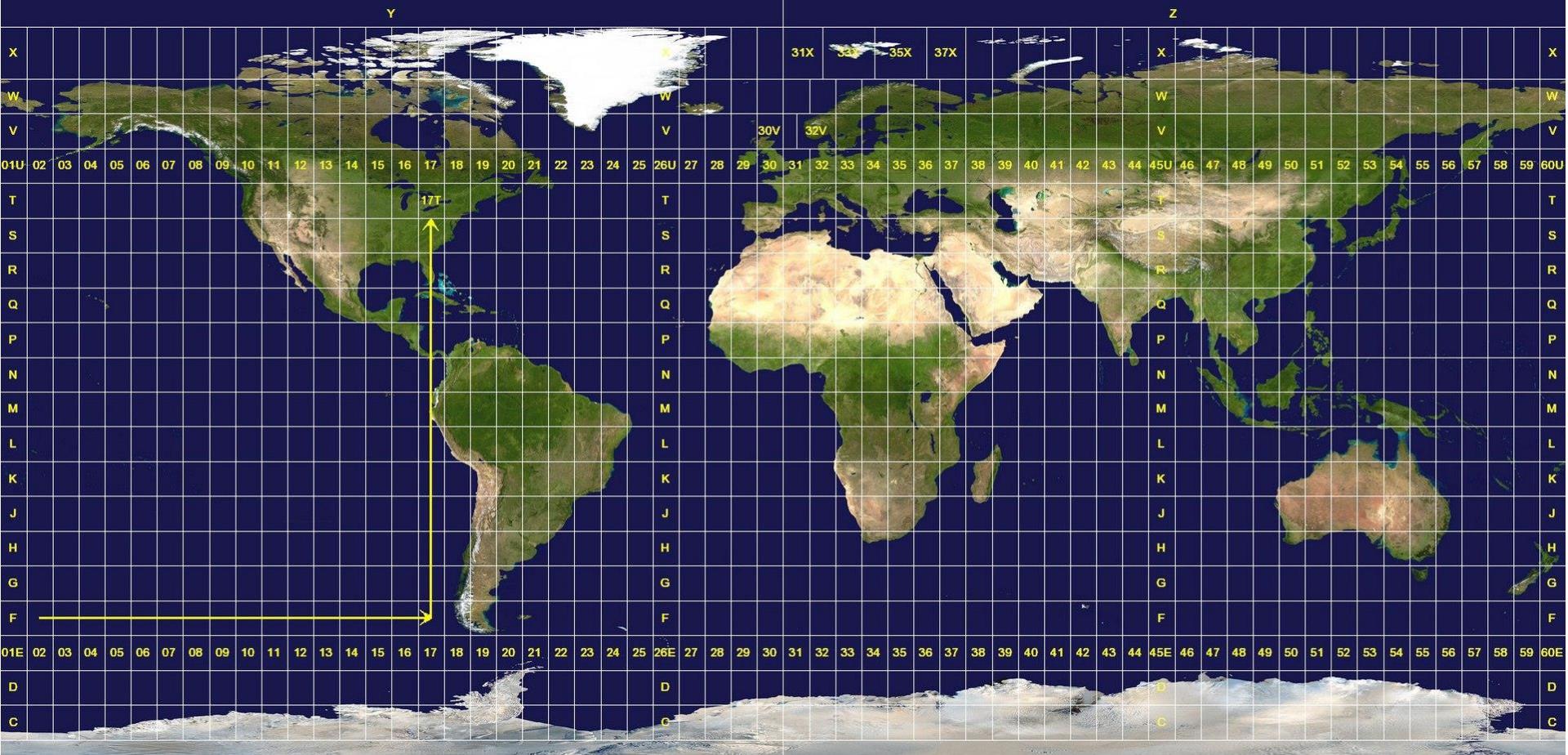
Source: Benning (2009)

- Zone is 6° wide
- Projection is isogonal / equal of angle
- Northing value of the equator: $N = 0 \text{ m}$
- Easting value of each central meridian: $E = 500'000 \text{ m}$
- Scale factor at central meridian: 0.9996
- Zone numbering commences with 1 in the zone 180°W to 174°W and increases eastward to zone 60 at the zone 174°E to 180°E
- Projection limits of latitude 80°S to 80°N



East [m]	32 364939
North [m]	5621299

- 60 stripes of 6° each = 360°



UTM projection: European grid



Fig.: UTM Zones in Europa
Source: www.wikipedia.org

- Germany: stripes 31, 32, 33
- North Rhine-Westphalia: only stripe 32 with central meridian at 9°

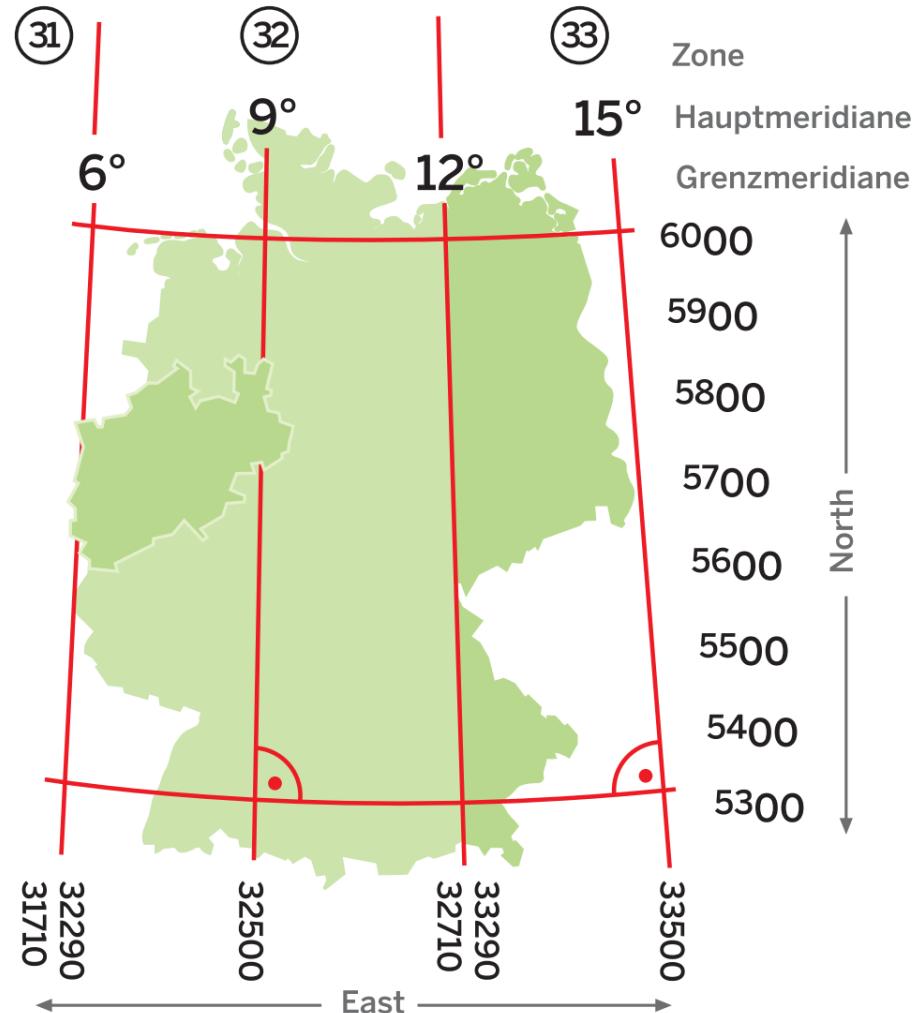


Abb. 6: Die Lage von NRW in der UTM-Zone 32

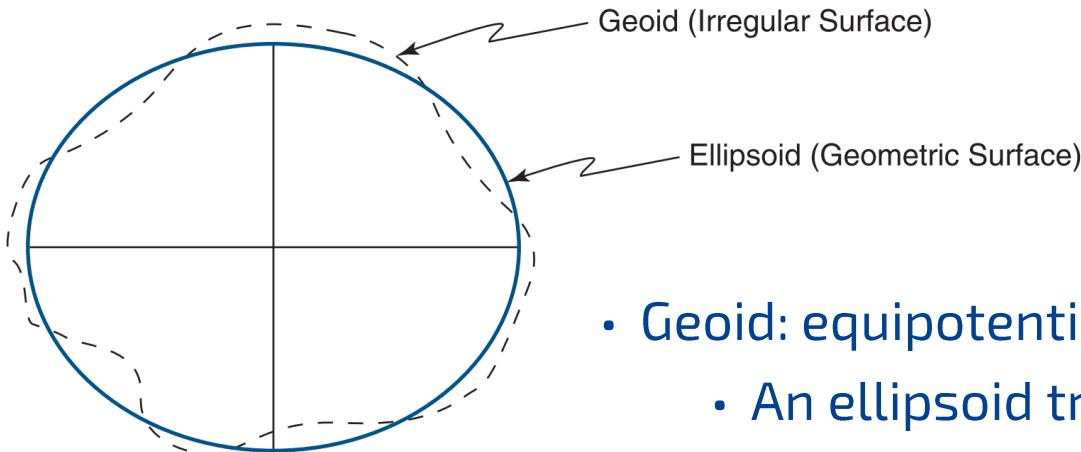
Source: Geobasis NRW

- UTM projection not directly related to any datum nor reference ellipsoid
- Countries may choose different datum and reference ellipsoids
- Germany:
 - Datum = ETRS89
 - Reference Ellipsoid: GRS80 Ellipsoid

	a	1/f
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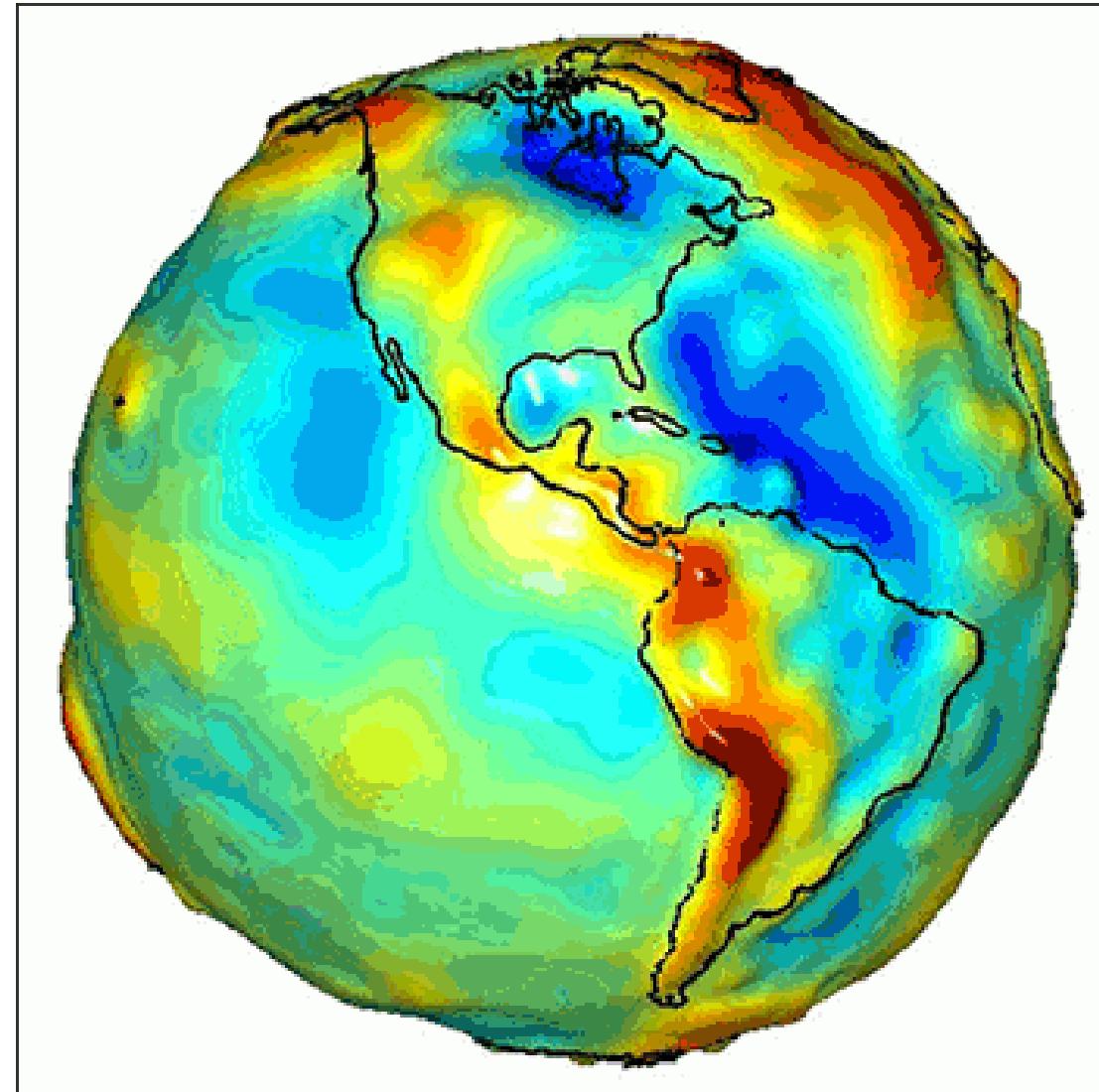
- Physical heights are needed
 - => no water flows between two points with the same height
 - => Equal to height above geoid



- Geoid: equipotential surface of the gravity field
 - An ellipsoid tries to approximate the geoid
 - Difference between geoid and ellipsoid = Geoid undulation
- => Geoid undulation must be known to get physical heights out of ellipsoidal heights

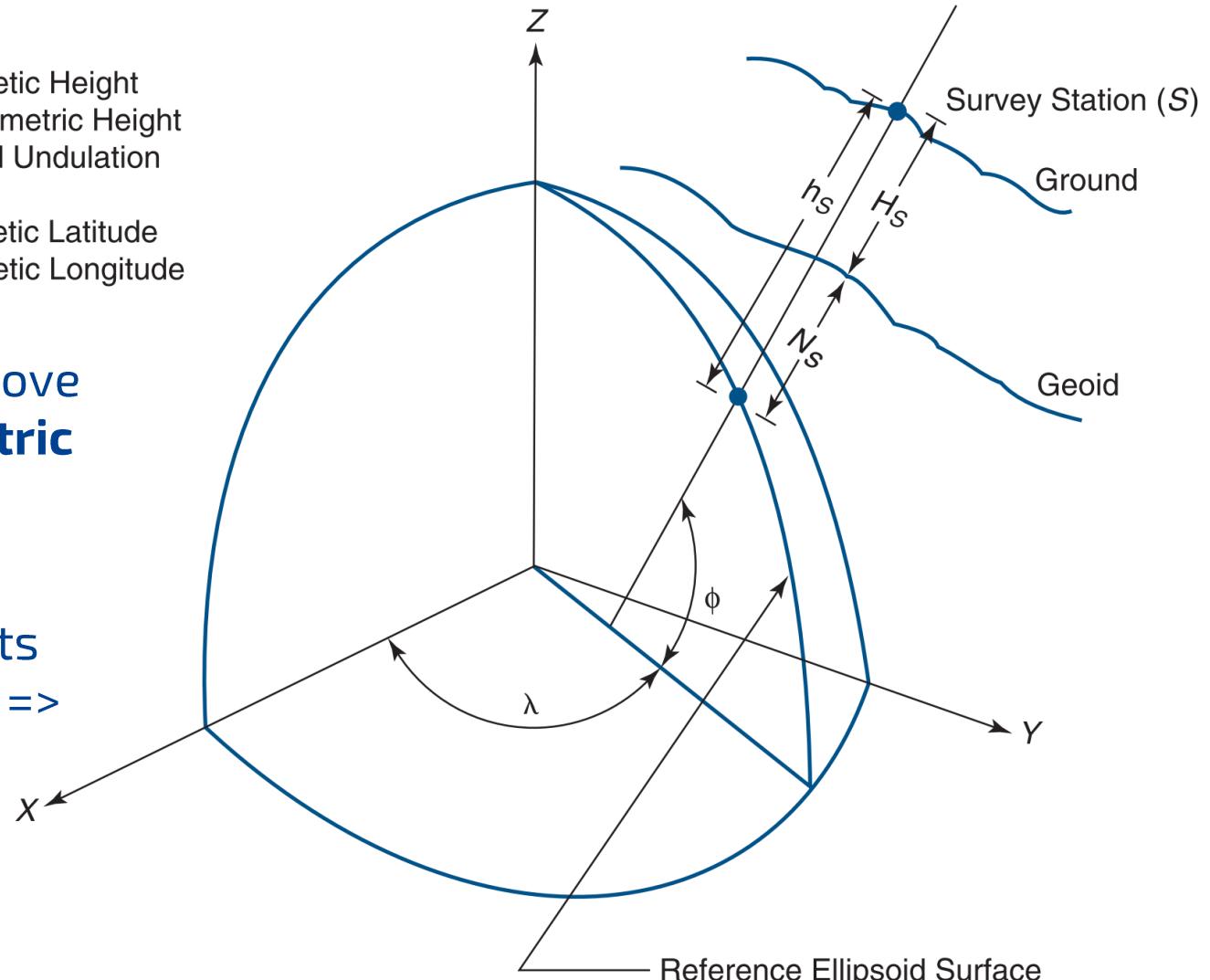
Source:
Kavanagh
(2014)

Physical heights



h = Geodetic Height
 H = Orthometric Height
 N = Geoid Undulation
 $h = H + N$
 ϕ = Geodetic Latitude
 λ = Geodetic Longitude

- Physical height above geoid => **orthometric heights**

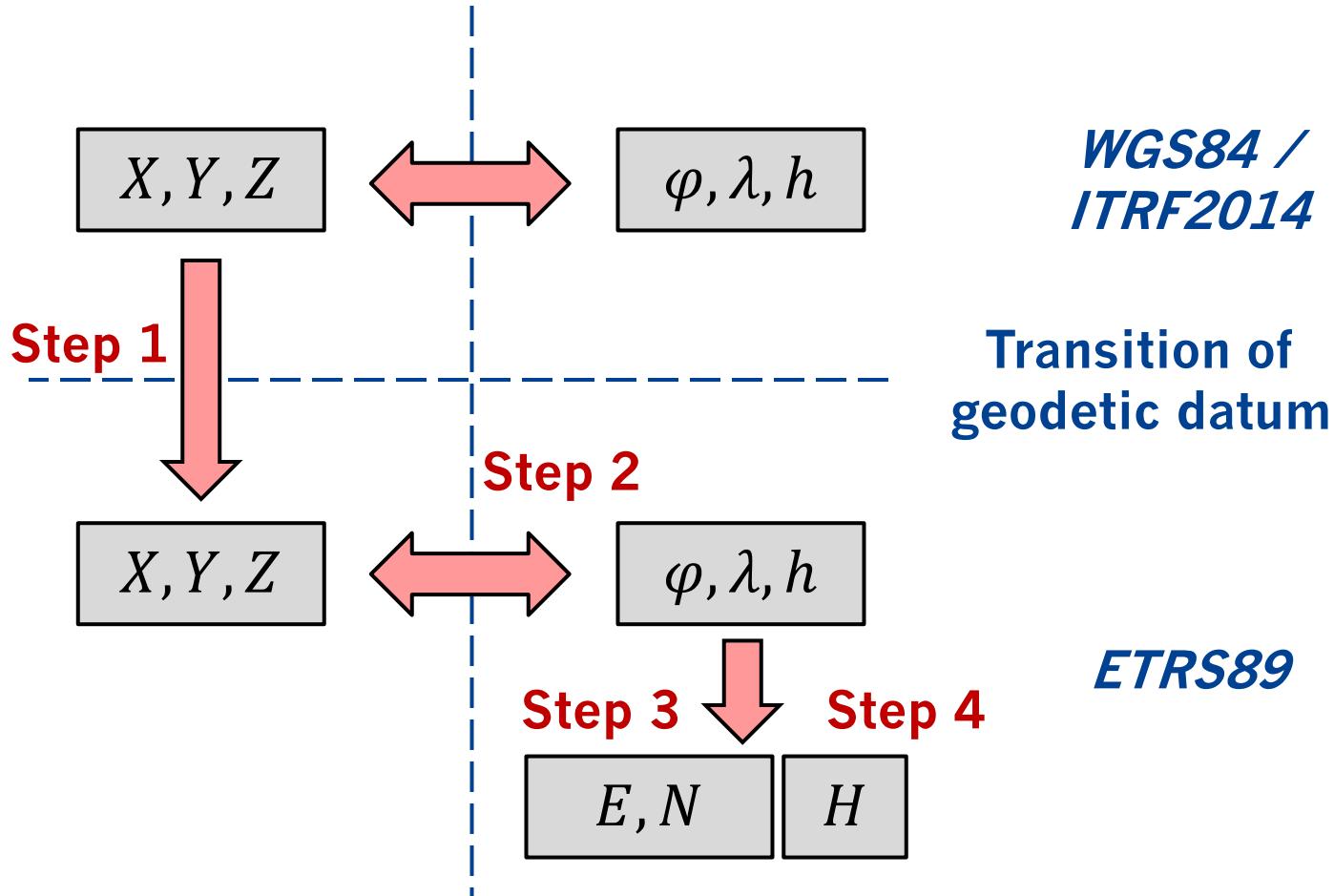


Source:
Kavanagh
(2014)

- **Positional coordinate**
 - UTM projection
 - Datum: ETRS89
 - Reference ellipsoid: GRS80
- **Height**
 - Physical heights: normal heights
 - Datum: GCG2016 (if transforming from ellipsoidal to physical)
 - Reference: Quasi geoid

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Transition of reference surface

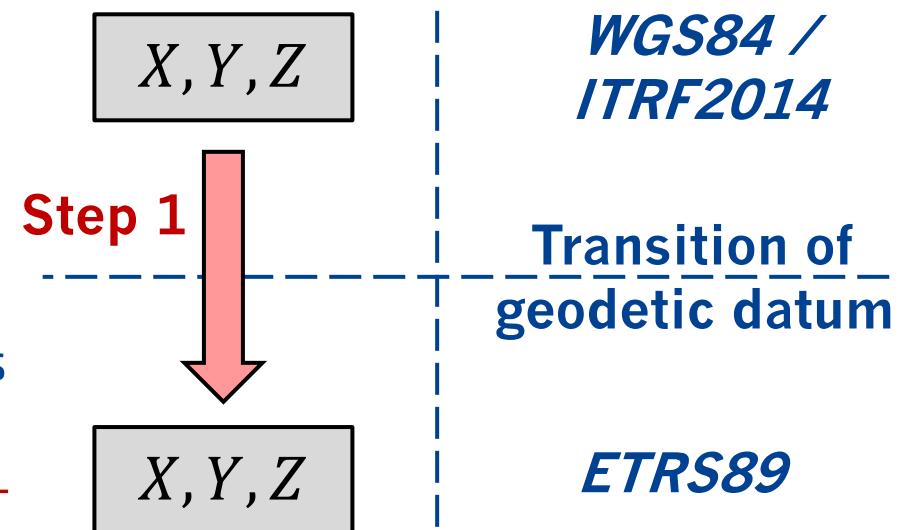


- Absolute positioning with GNSS:
 - Absolute coordinates X, Y, Z are given in WGS84
 - Transformation parameters between WGS84 and ETRS89 are frequently sent to GNSS receiver

=> Datum transition by selection in receiver => **X, Y, Z in ETRS89**
- Relative positioning with GNSS:
 - Baselines $\Delta X, \Delta Y, \Delta Z$ are given in WGS84
 - Similarity transformation between WGS84 and ETRS89
approx. pure translation

=> baselines unaffected from datum transition

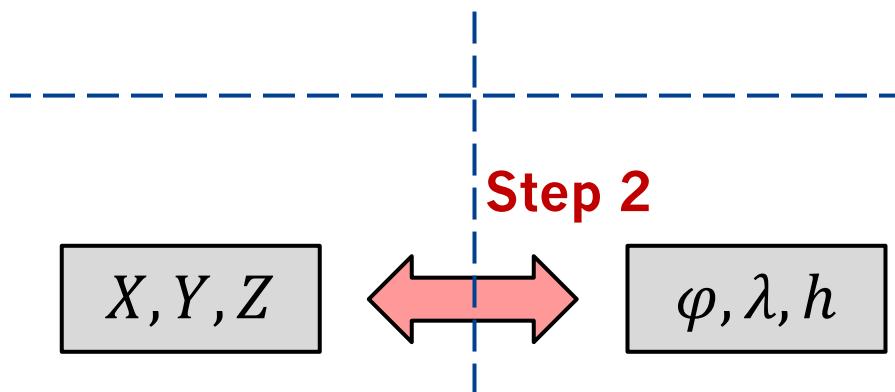
=> Use master station coordinates with datum ETRS89 and add baselines $\Delta X, \Delta Y, \Delta Z$ => **$X + \Delta X, Y + \Delta Y, Z + \Delta Z$ in ETRS89**



Step 2: From X, Y, Z to φ, λ, h

$$\begin{bmatrix} \varphi \\ \lambda \\ h \end{bmatrix} = \begin{bmatrix} \tan^{-1} \left(\frac{z}{\sqrt{x^2+y^2}} \cdot \left(1 - \frac{a^2-b^2}{a^2} \cdot \frac{N(\varphi)}{N(\varphi)+h} \right)^{-1} \right) \\ \tan^{-1} \left(\frac{y}{x} \right) \\ \frac{\sqrt{x^2+y^2}}{\cos \varphi} - N(\varphi) \end{bmatrix}$$

- Iterate: $N(\varphi) = \frac{a^2}{\sqrt{a^2 \cdot \cos^2 \varphi + b^2 \cdot \sin^2 \varphi}}$
- Use GRS80 Ellipsoid



Transition of
geodetic datum

ETRS89

- $E = E_0 + [1] \cdot \Delta L + [3] \cdot \Delta L^3 + [5] \cdot \Delta L^5$
- $N = m \cdot G + [2] \cdot \Delta L^2 + [4] \cdot \Delta L^4 + [6] \cdot \Delta L^6$
- $\Delta L = L - L_o$, with $L_o = 9^\circ$ for North Rhine-Westfalia
- $E_0 = \left(\frac{L_o + 3^\circ}{6^\circ} + 30.5 \right) \cdot 10^6$
- $m = 0.9996$
- $G = G_o \cdot B + G_2 \cdot \sin 2B + G_4 \cdot \sin 4B + G_6 \cdot \sin 6B$
- $[1] = \frac{m}{\rho} \cdot \bar{N} \cdot \cos B$
- $[3] = \frac{m}{6\rho^3} \cdot \bar{N} \cdot \cos^3 B \cdot (1 - t^2 + \eta^2)$
- $[5] = \frac{m}{120\rho^5} \cdot \bar{N} \cdot \cos^5 B \cdot (5 - 18t^2 + t^4 + \eta^2 \cdot (14 - 58t^2))$
- $[2] = \frac{m}{2\rho^2} \cdot \bar{N} \cdot \cos^2 B \cdot t$
- $[4] = \frac{m}{24\rho^4} \cdot \bar{N} \cdot \cos^4 B \cdot t \cdot (5 - t^2 + 9 \cdot \eta^2)$
- $[6] = \frac{m}{720\rho^6} \cdot \bar{N} \cdot \cos^6 B \cdot t \cdot (61 - 58t^2 + t^4)$

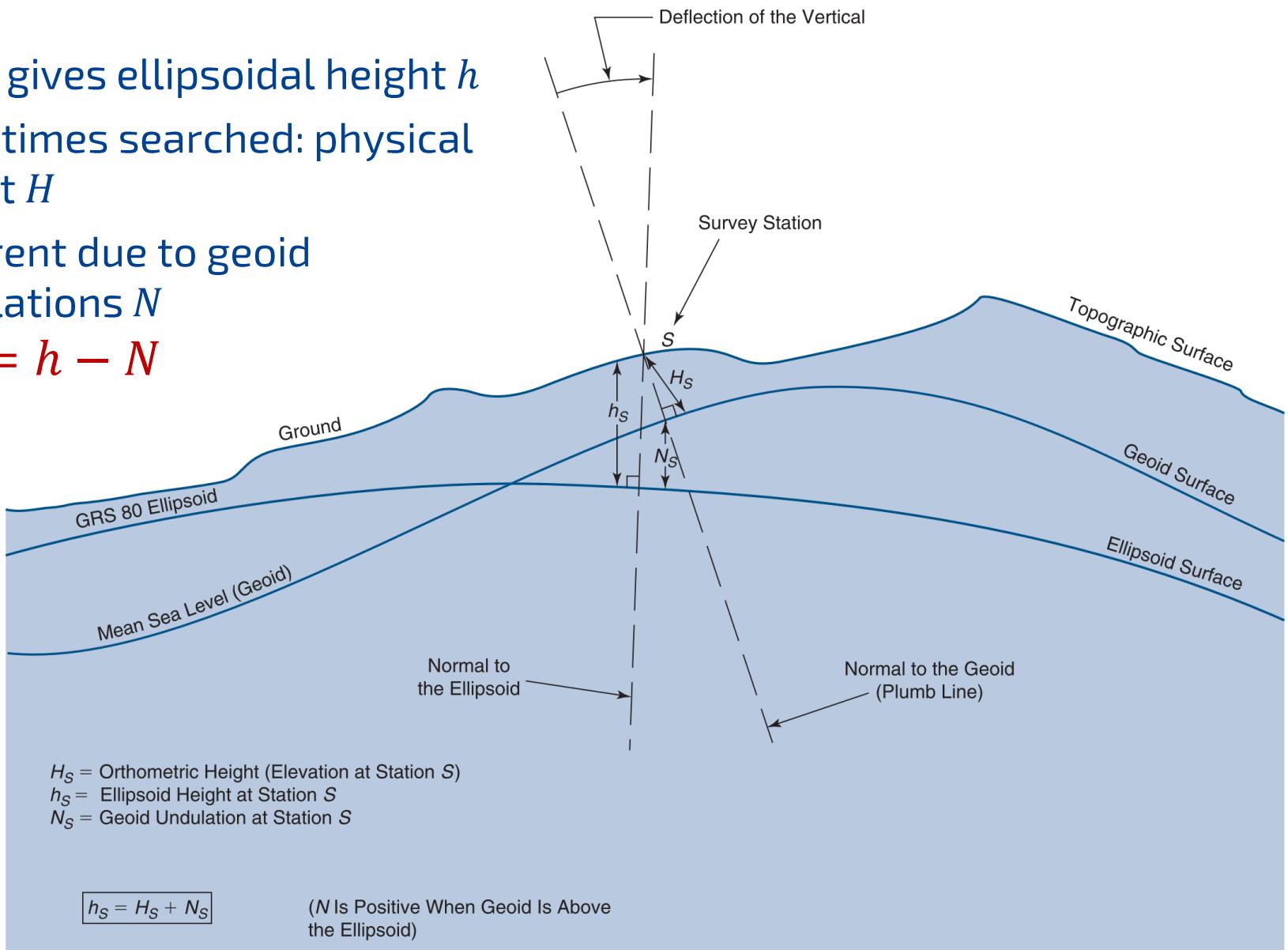
- $\rho = \frac{180^\circ}{\pi}$
- $\bar{N} = \frac{c}{\sqrt{1+\eta^2}}$
- $t = \tan B$
- $\eta^2 = \frac{a^2-b^2}{b^2} \cdot \cos^2 B$
- $c = \frac{a^2}{b}$

$\lambda = L = \text{Longitude}$
 $\varphi = B = \text{Latitude}$

Parameter	GRS80 ellipsoid
G_o	$111'132.952547 \text{ m/}^\circ$
G_2	$-16'038.5088 \text{ m}$
G_4	16.8326 m
G_6	-0.0220 m
a	$6'378'137.000 \text{ m}$
b	$6'356'752.314 \text{ m}$

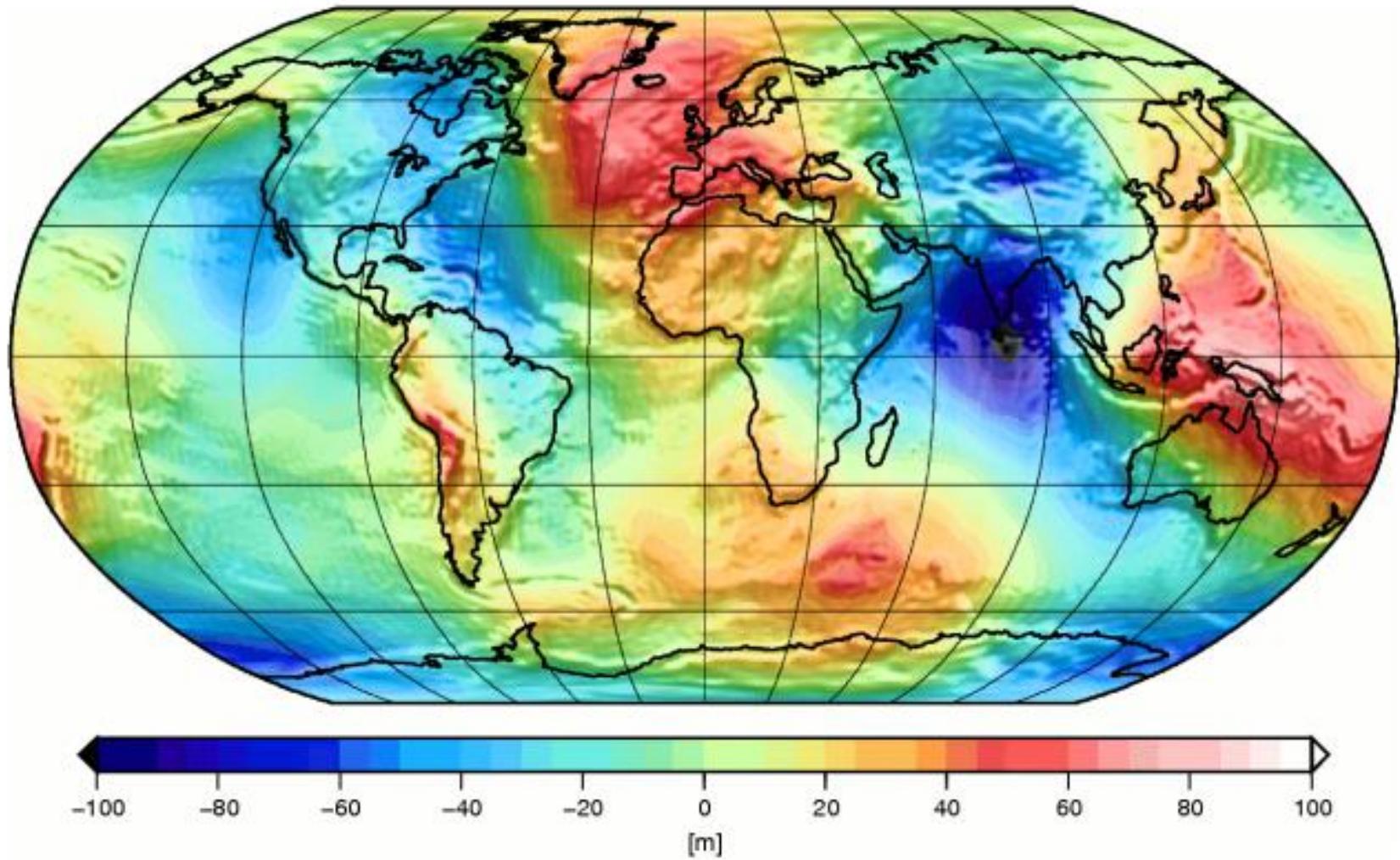
Step 4: From h to H

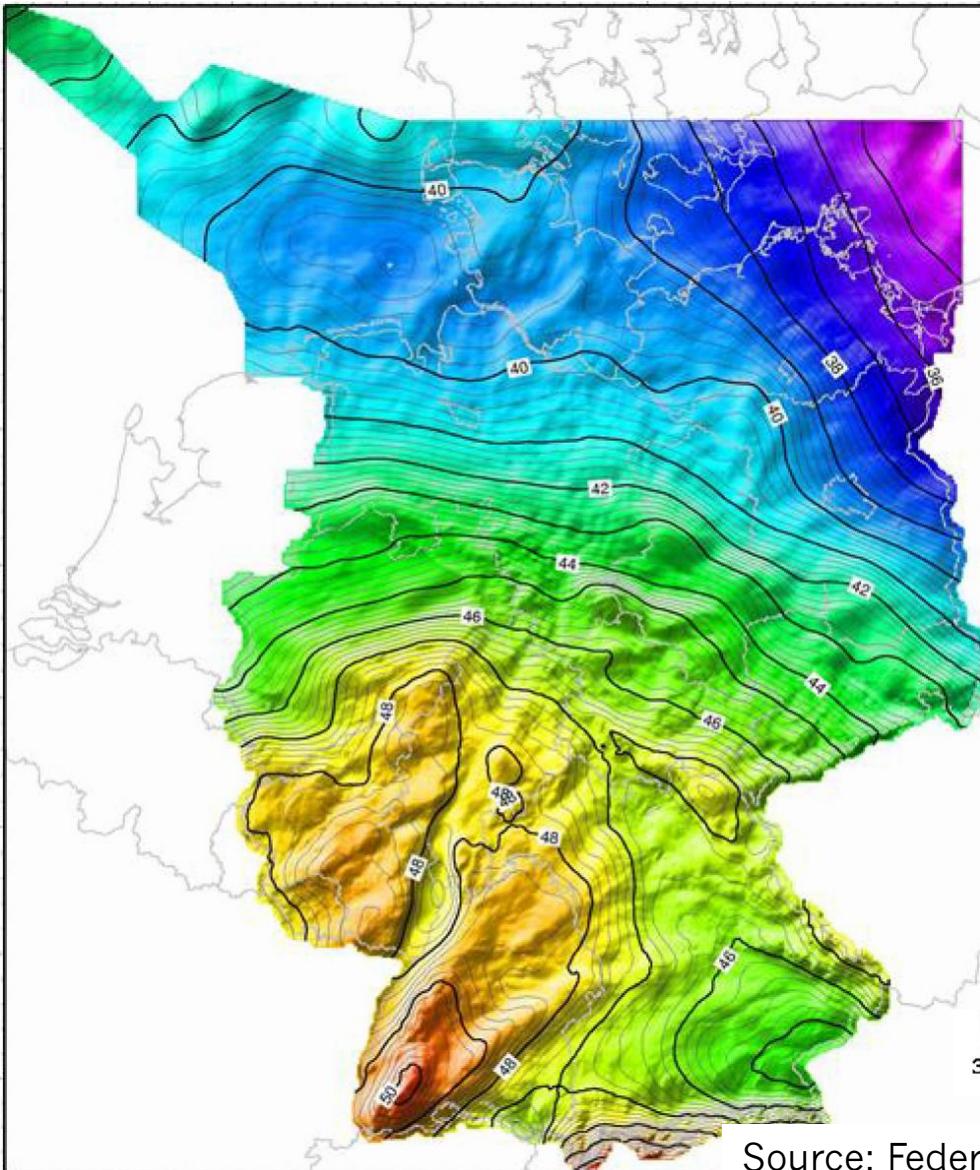
- GNSS gives ellipsoidal height h
 - Most times searched: physical height H
 - Different due to geoid undulations N
- $\Rightarrow H = h - N$



Source:
Kavanagh
(2014)

Geoid undulation





- Federal Office for Cartography and Geodesy provides quasi geoid undulations for Germany
- Resolution: approx. 1 km x 1 km.
- Accuracy: 2-3 cm
- North Rhine-Westphalia: 42 m ... 48 m
- Germany: 34 m ... 50 m

Source: Federal Office for Cartography and Geodesy, Germany

- Step 1: Transform GNSS coordinate with datum WGS84 to datum ETRS89 (made by GNSS receiver that knows frequently updated transformation parameters)

=> Transition of geodetic datum

- Step 2: Transform Cartesian coordinates to ellipsoidal coordinates using GRS80 ellipsoid

=> Transition of reference surface

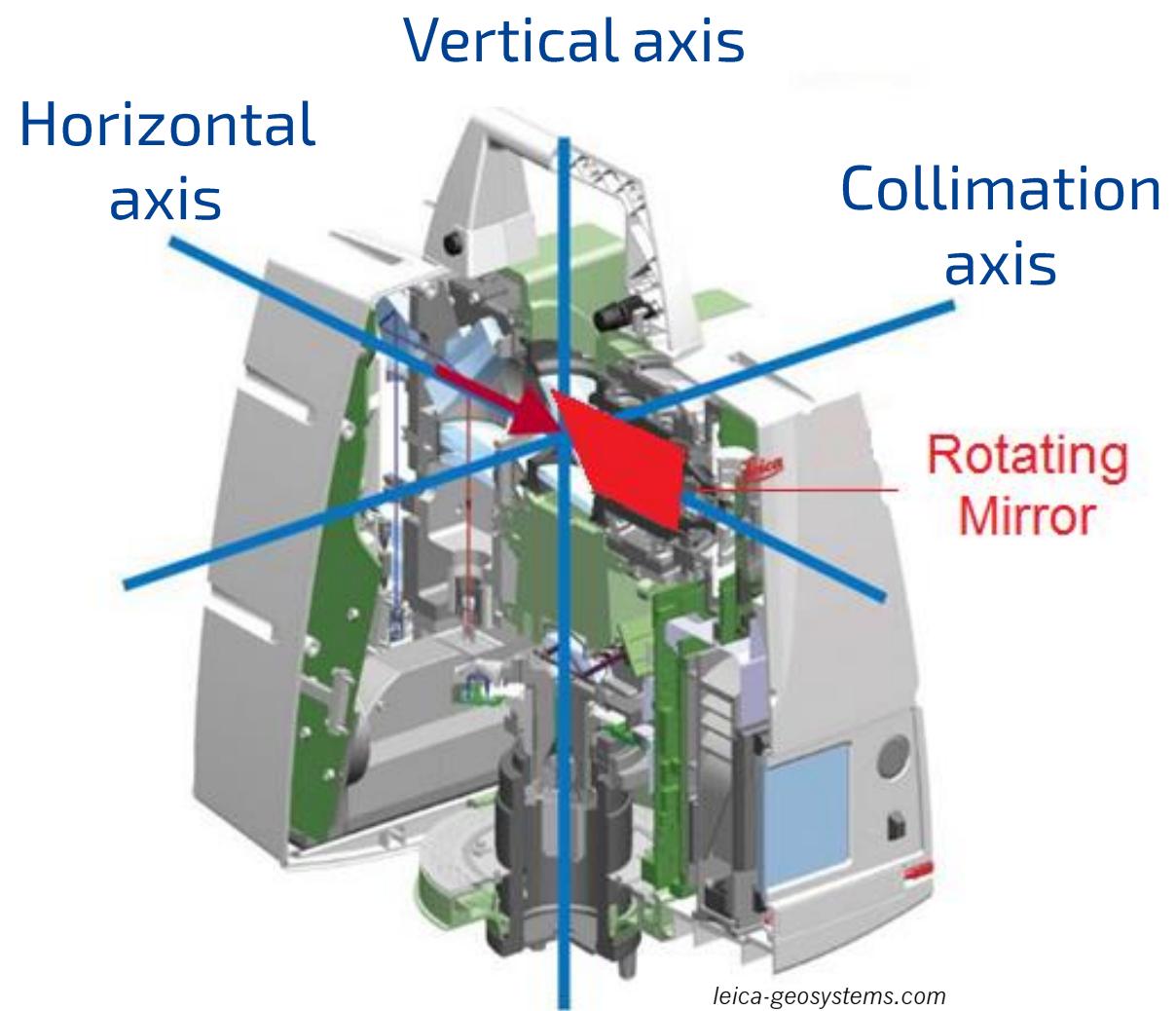
- Step 3: Transform latitude and longitude to UTM coordinates

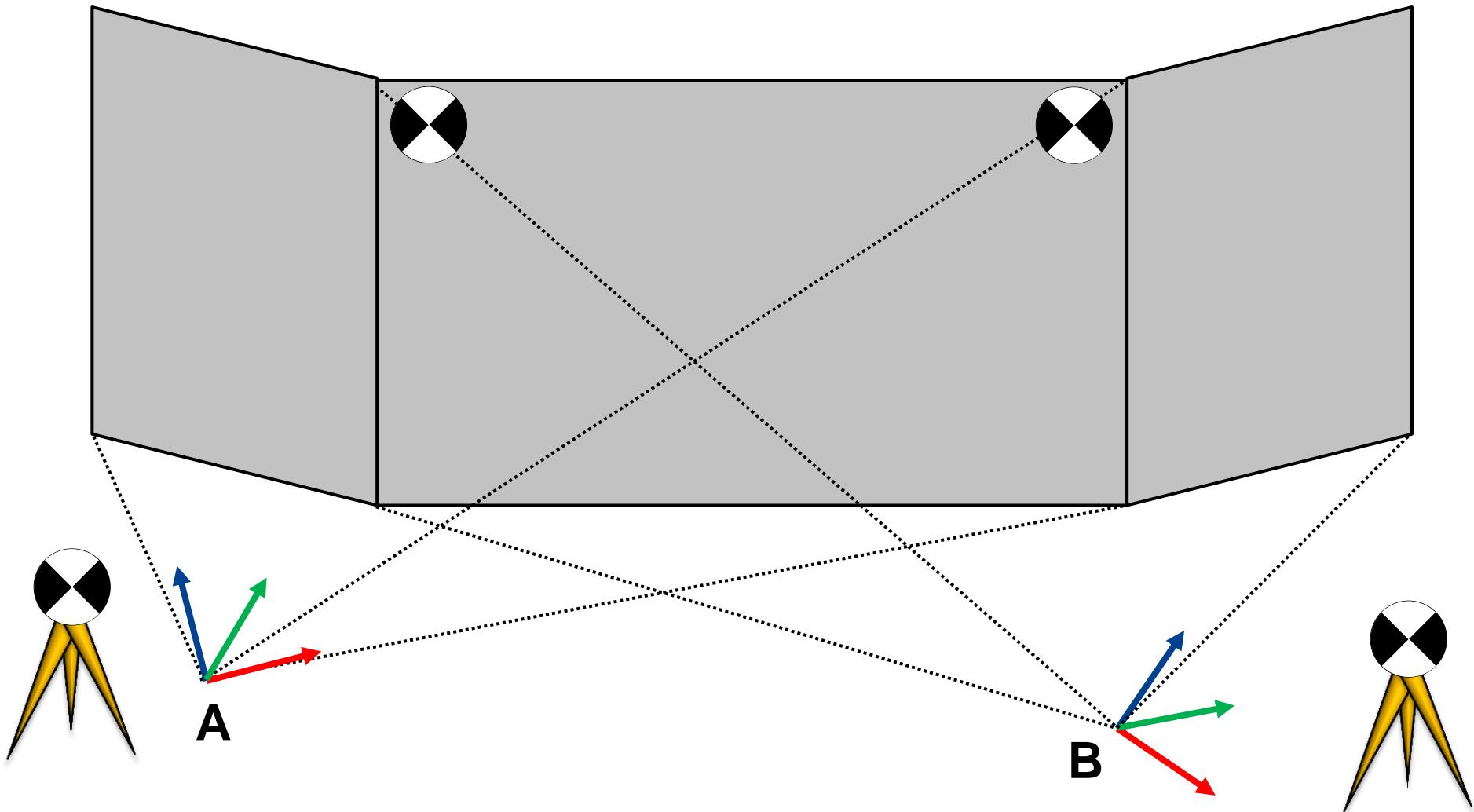
=> Projection

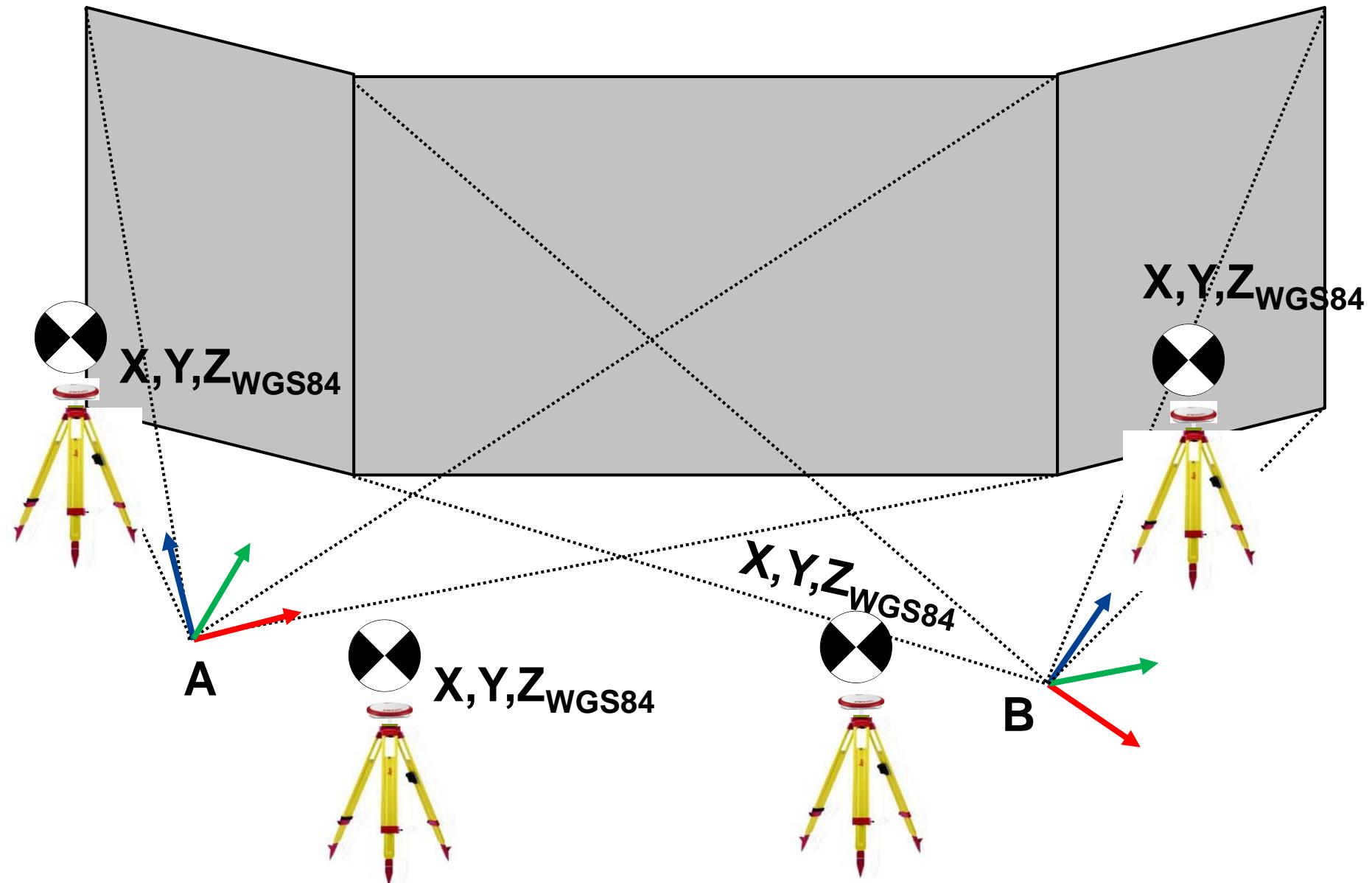
- Step 4: Transform ellipsoidal height to physical height using the GCG2016 quasi geoid undulations

=> Transition of reference surface

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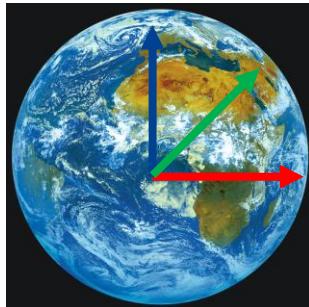






Process each single laser scan point similarly to GNSS coordinates;

Start from Step 1 (WGS84) or Step 2 (ETRS89)

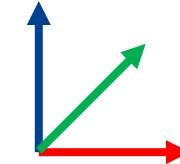


X, Y, Z_{WGS84} or
 X, Y, Z_{ETR89} (absolute positions)

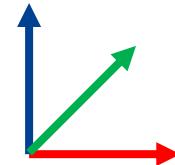
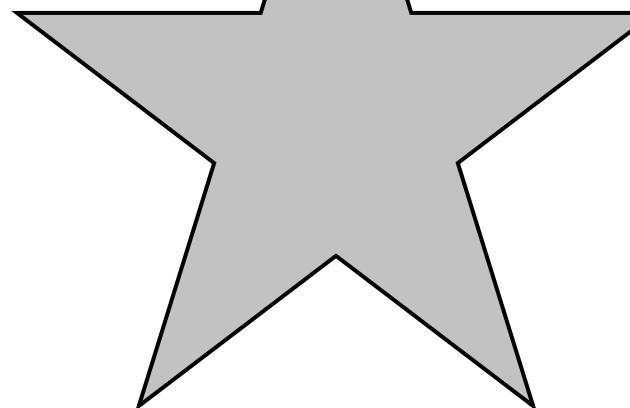


$t_{x,1}, t_{y,1},$
 $t_{z,1}, \varepsilon_{z,1}$

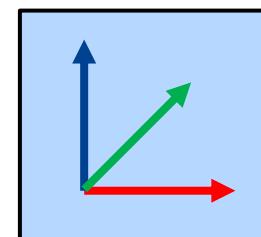
$t_{x,2}, t_{y,2},$
 $t_{z,2}, \varepsilon_{z,2}$



$t_{x,3}, t_{y,3},$
 $t_{z,3}, \varepsilon_{z,3}$



$t_{x,4}, t_{y,4},$
 $t_{z,4}, \varepsilon_{z,4}$



$t_X, t_Y, t_Z,$
 $\varepsilon_X, \varepsilon_Y, \varepsilon_Z$

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